



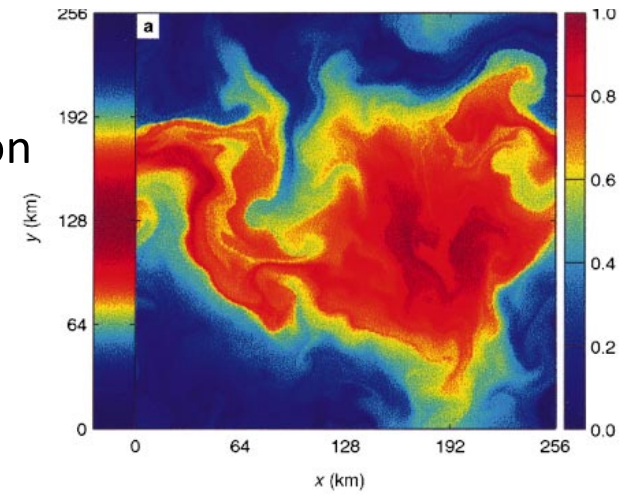
A satellite image of Earth showing the Western Hemisphere, including North and South America. The oceans are color-coded to represent chlorophyll-a concentrations, with green and cyan indicating higher concentrations and blue indicating lower concentrations. The text "Biogeochemical proxies from optics" is overlaid in a semi-transparent dark box across the center of the image.

# Biogeochemical proxies from optics

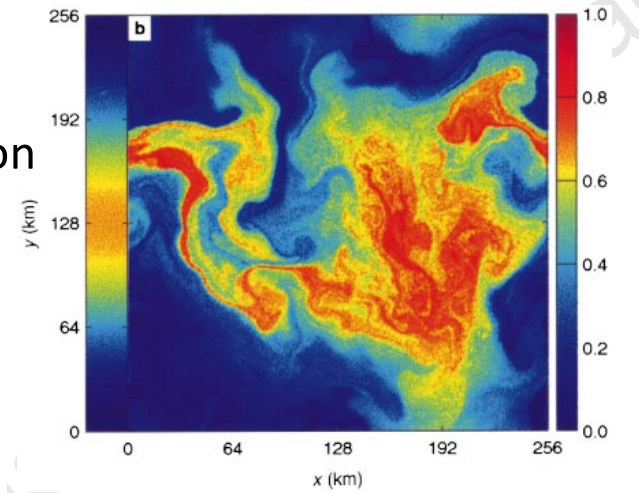
Meg Estapa, UMaine

Slides borrowed heavily from Ivona Cetinić, NASA GSFC

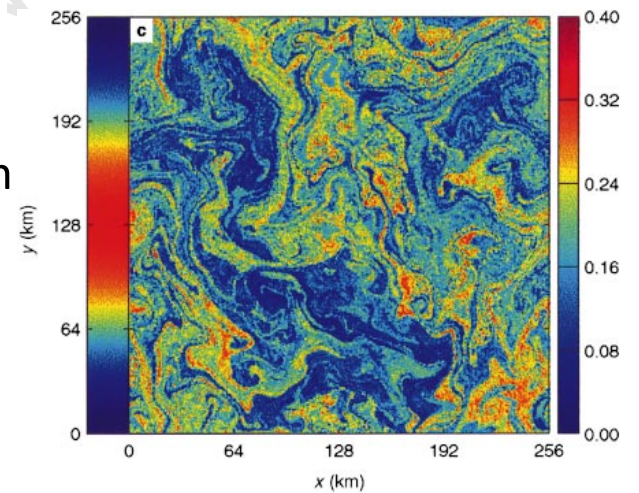
Phytoplankton  
“carrying  
capacity”



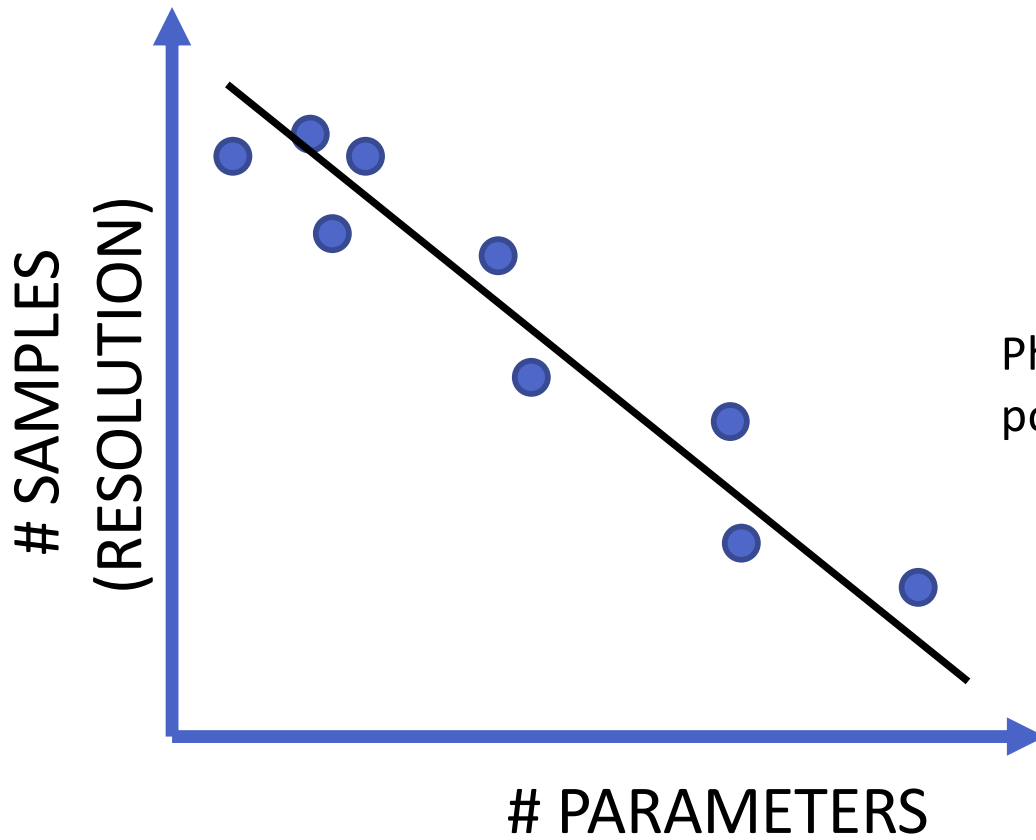
Phytoplankton  
population



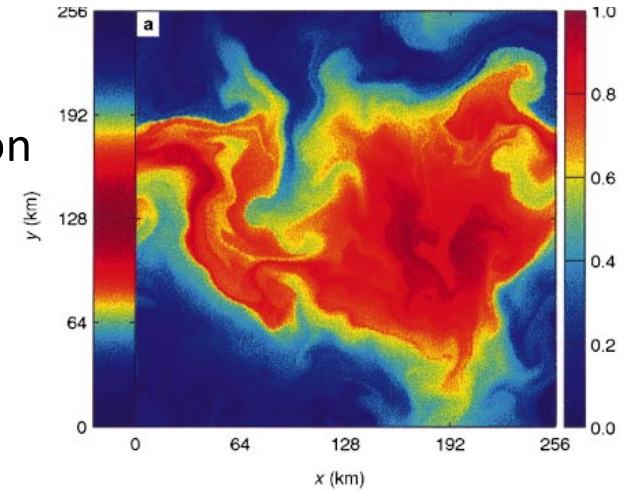
Zooplankton  
population



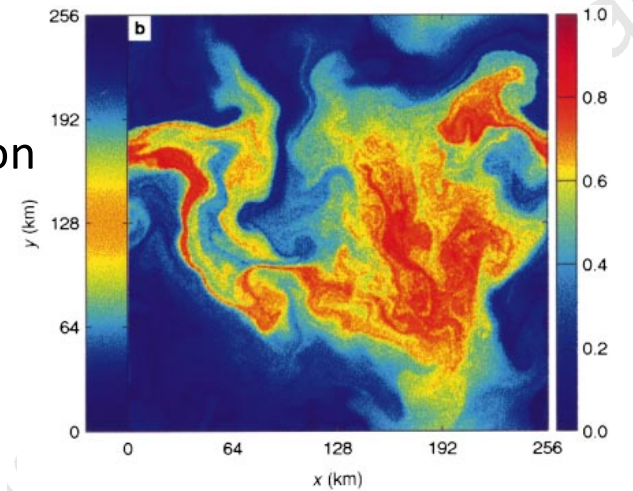
Abraham, 1998



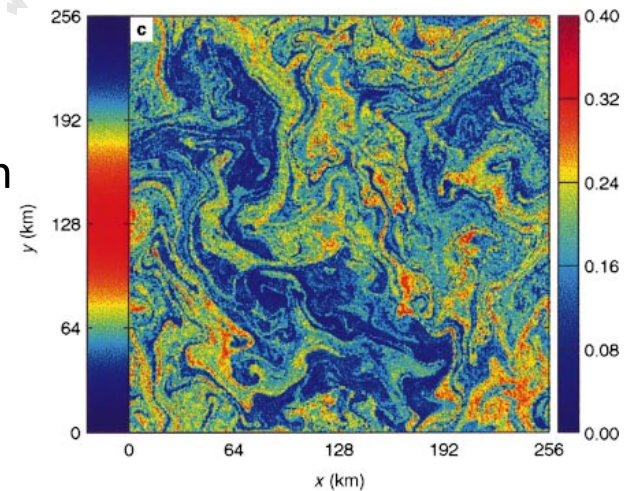
Phytoplankton  
“carrying capacity”



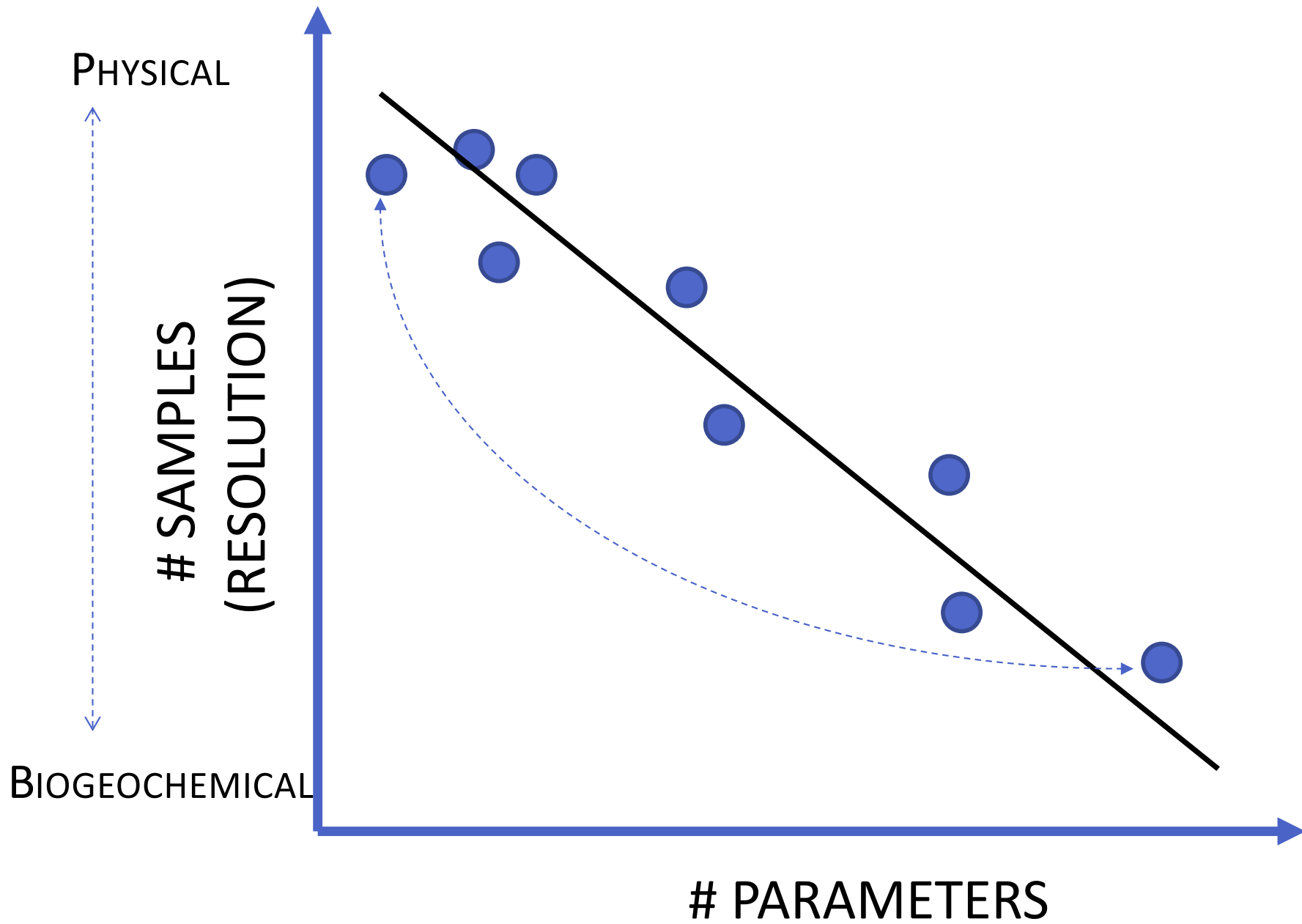
Phytoplankton  
population



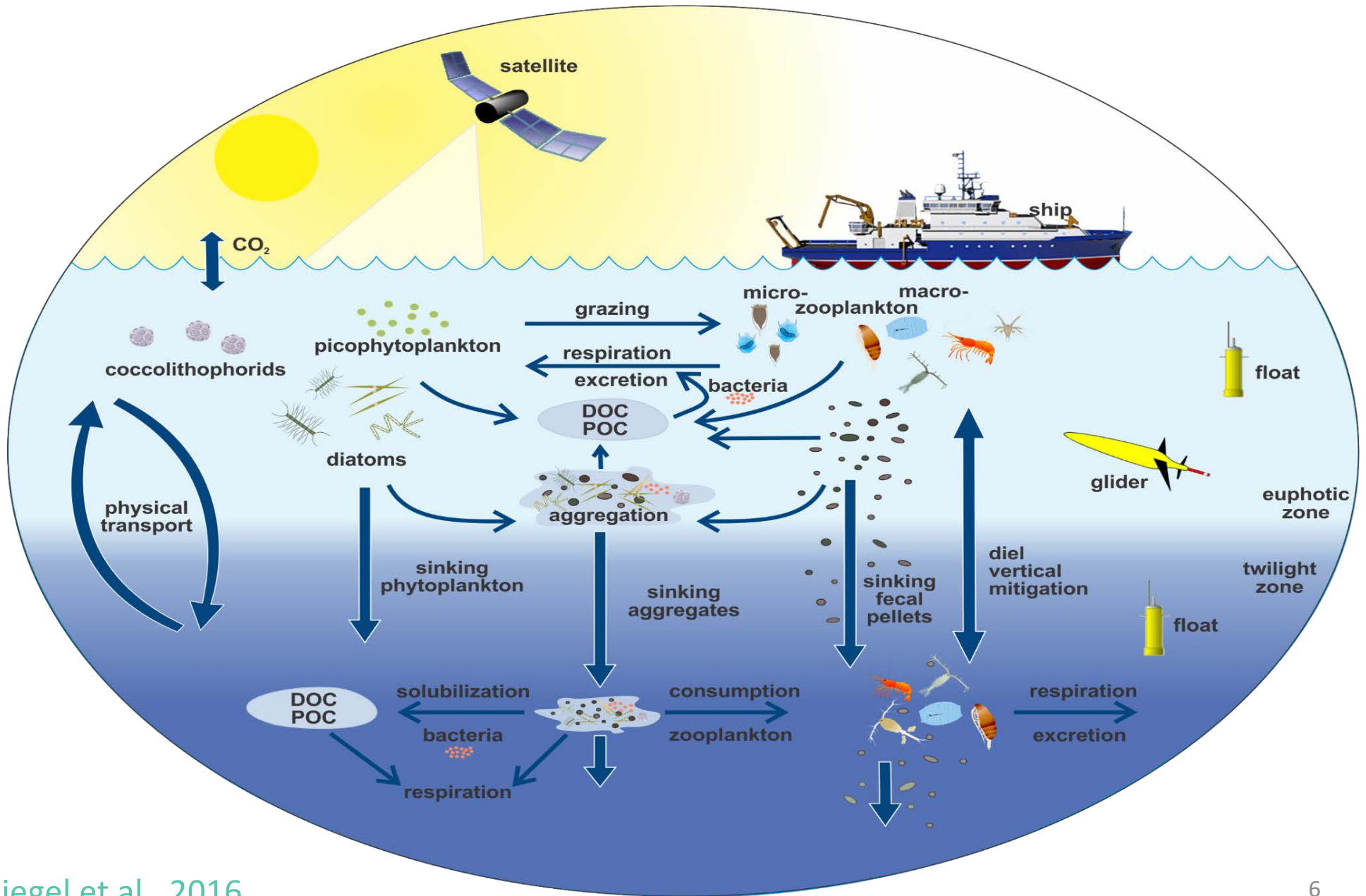
Zooplankton  
population



Abraham, 1998



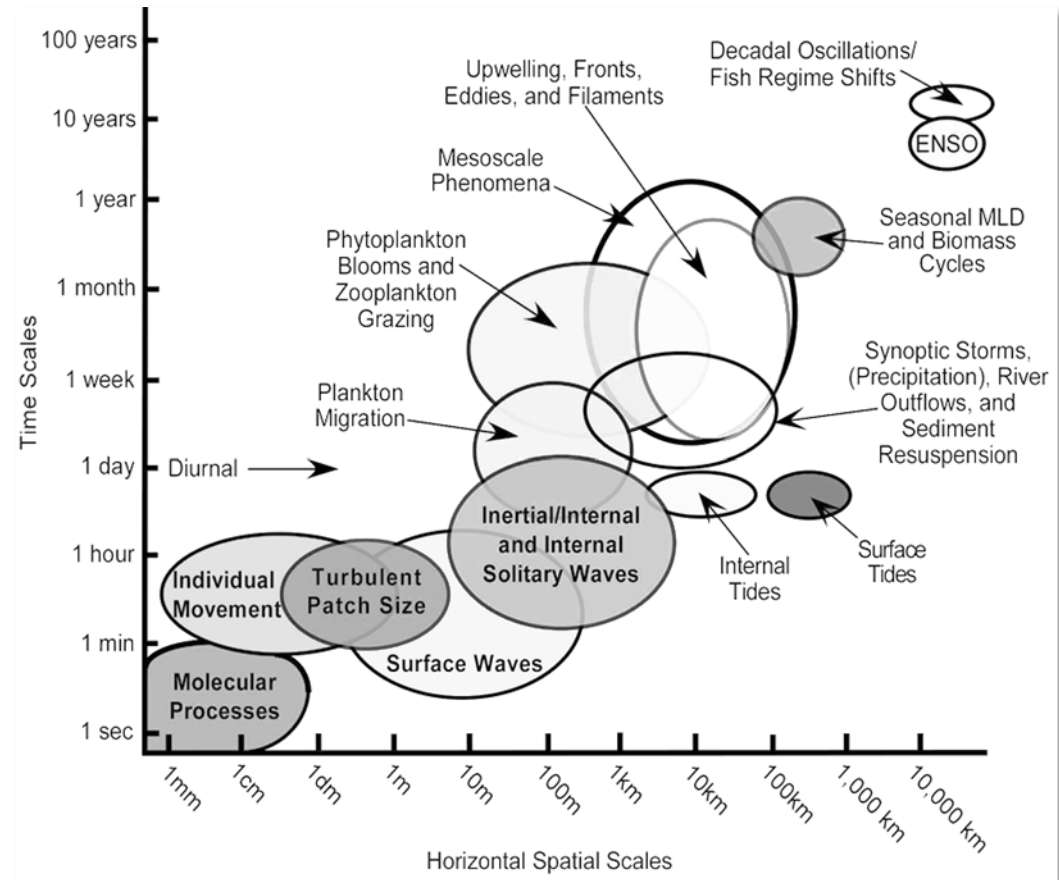
# Proxies are at the heart of major sampling programs



# Why?



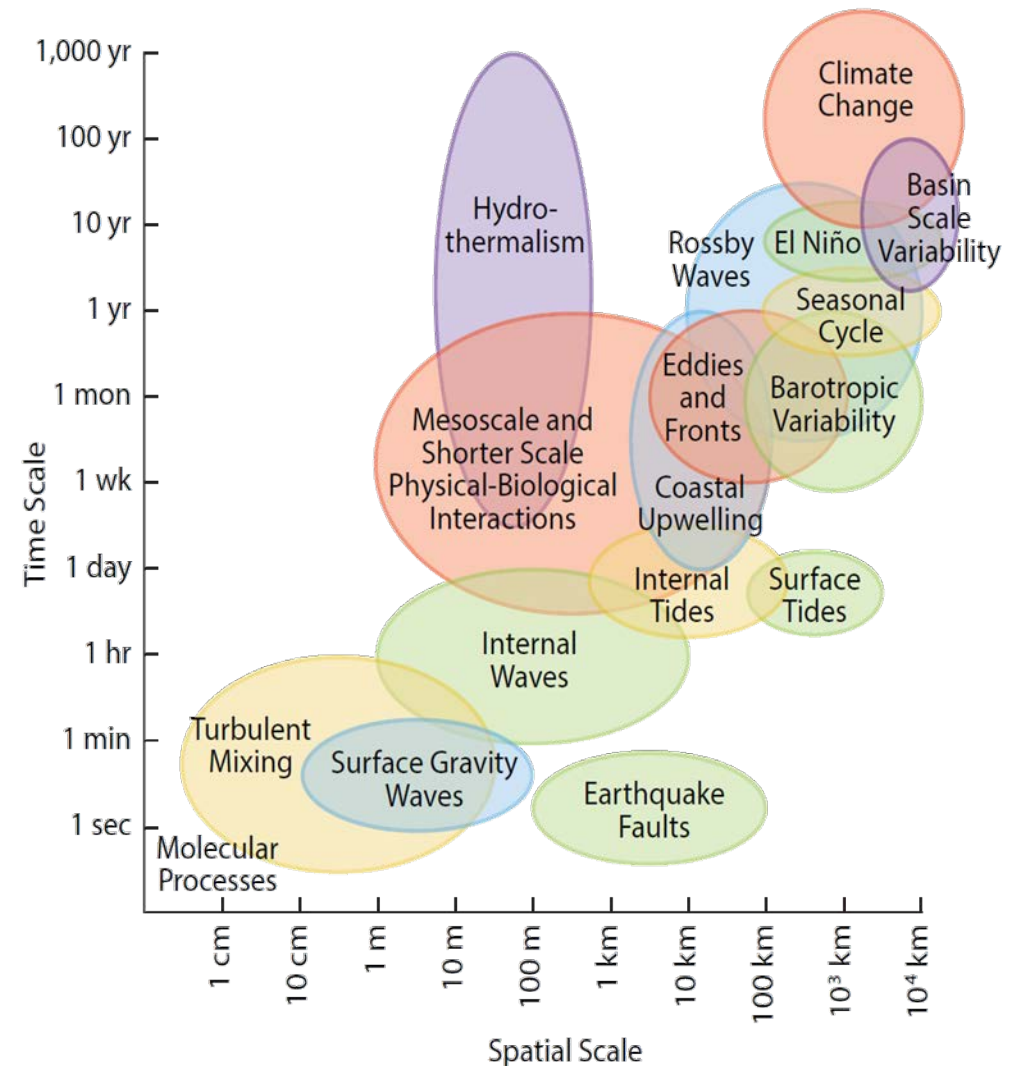
- Optics – in situ or remote sensed gives us higher resolution dataset
- Traditional methods (discrete) often expensive and time consuming
- Sampling the parameters on the scales of importance
- Validation for remote sensing and hi-res biogeochemical models (e.g. Haëntjens et al, 2017)



[Chang, G. and T. Dickey \(2008\).](#)

# Why?

- Optics – in situ or remote sensed gives us higher resolution dataset
- Traditional methods (discrete) often expensive and time consuming
- Sampling the parameters on the scales of importance
- Validation for remote sensing and hi-res biogeochemical models (e.g. Haëntjens et al, 2017)

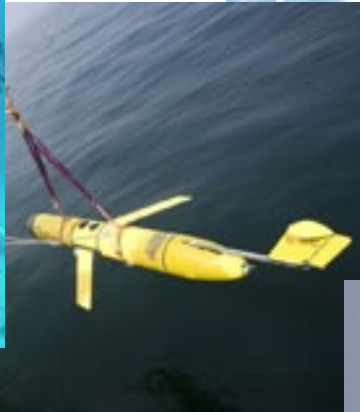


[Zykov and Miller \(2019\).](#)



# Why?

- Optical instruments are getting smaller, more robust and diverse
- They can be deployed over extended periods of time and in hard to reach areas



What are some *in situ* optical proxies (and associated biogeochemical quality/quantity)?

# Overview

- ✓ Where proxies fit, why to construct them
- Concentration proxies
  - “To first order, concentration controls IOPs”  
Emmanuel Boss, yesterday
- Composition and rate proxies
- Caveats

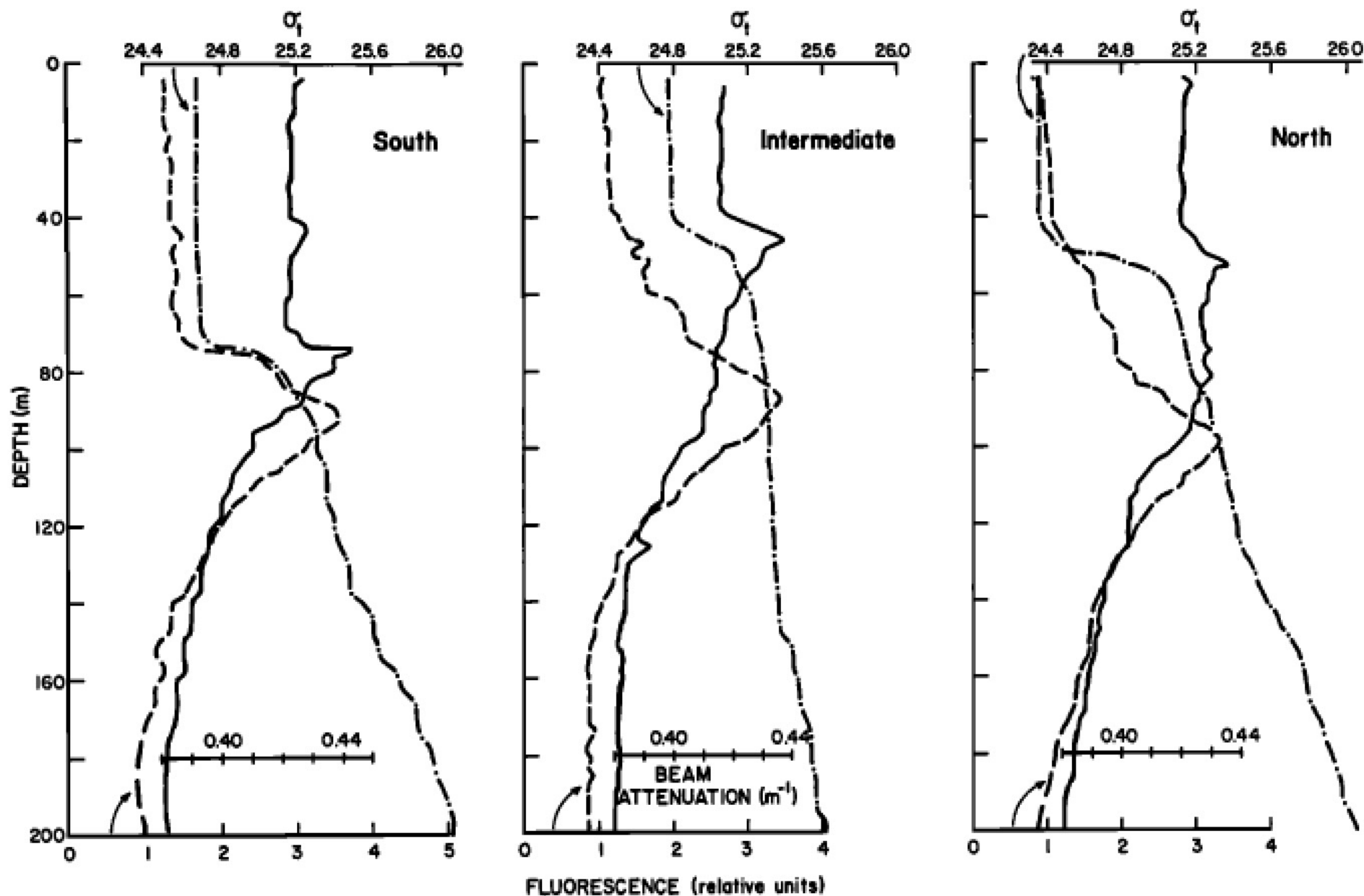
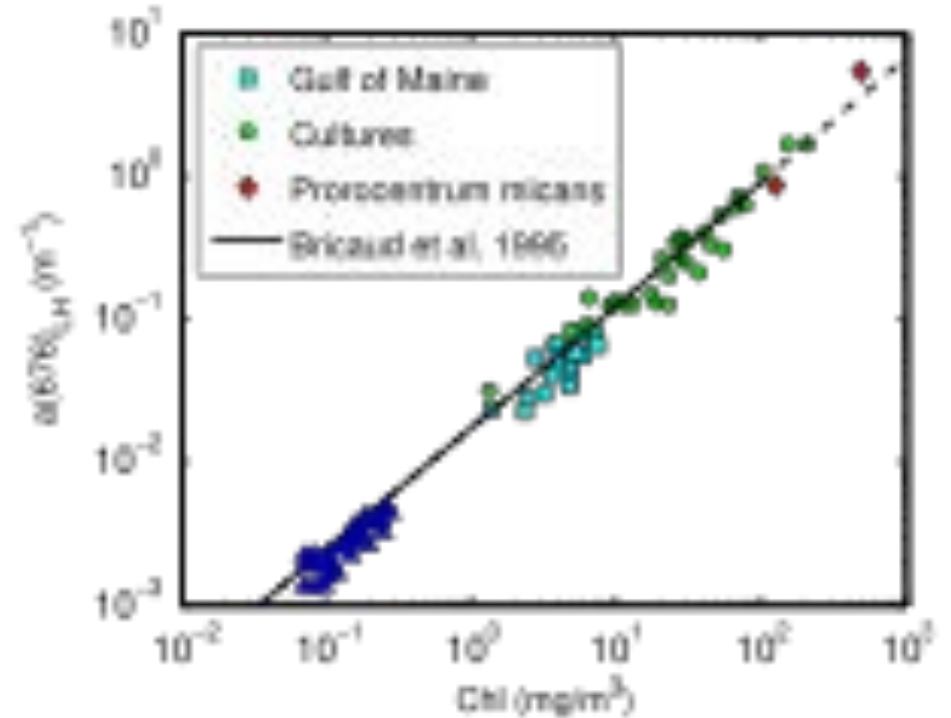
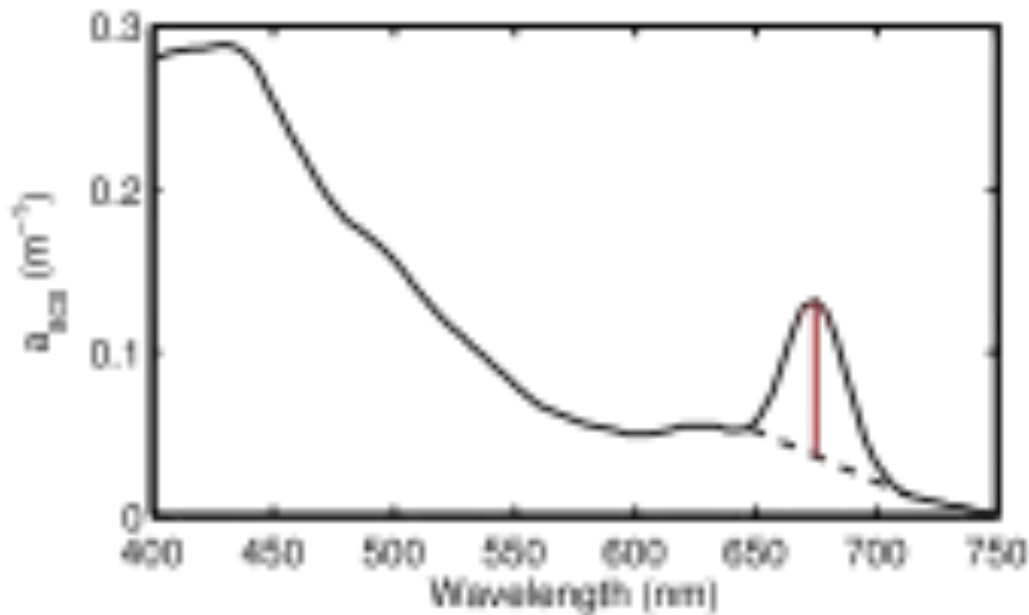


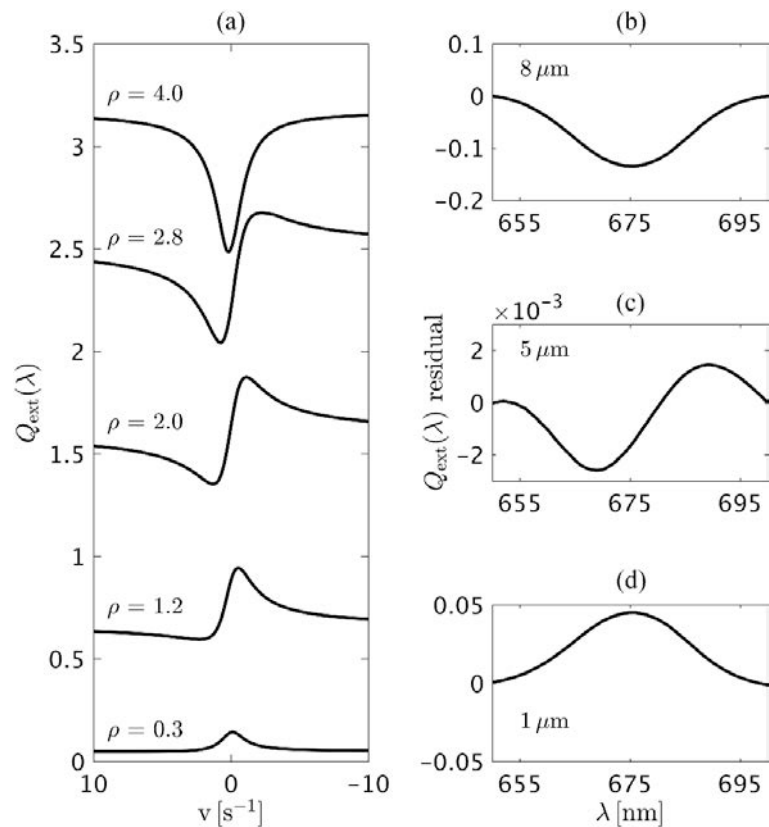
Fig. 1. Profiles of fluorescence, beam attenuation (665 nm) and  $\sigma_t$  for Pacific Central Gyre stations typical of waters north, south and in the Subtropical Front, Oct.-Nov., 1982.

# Absorption Line Height $\rightarrow$ Chlorophyll proxy

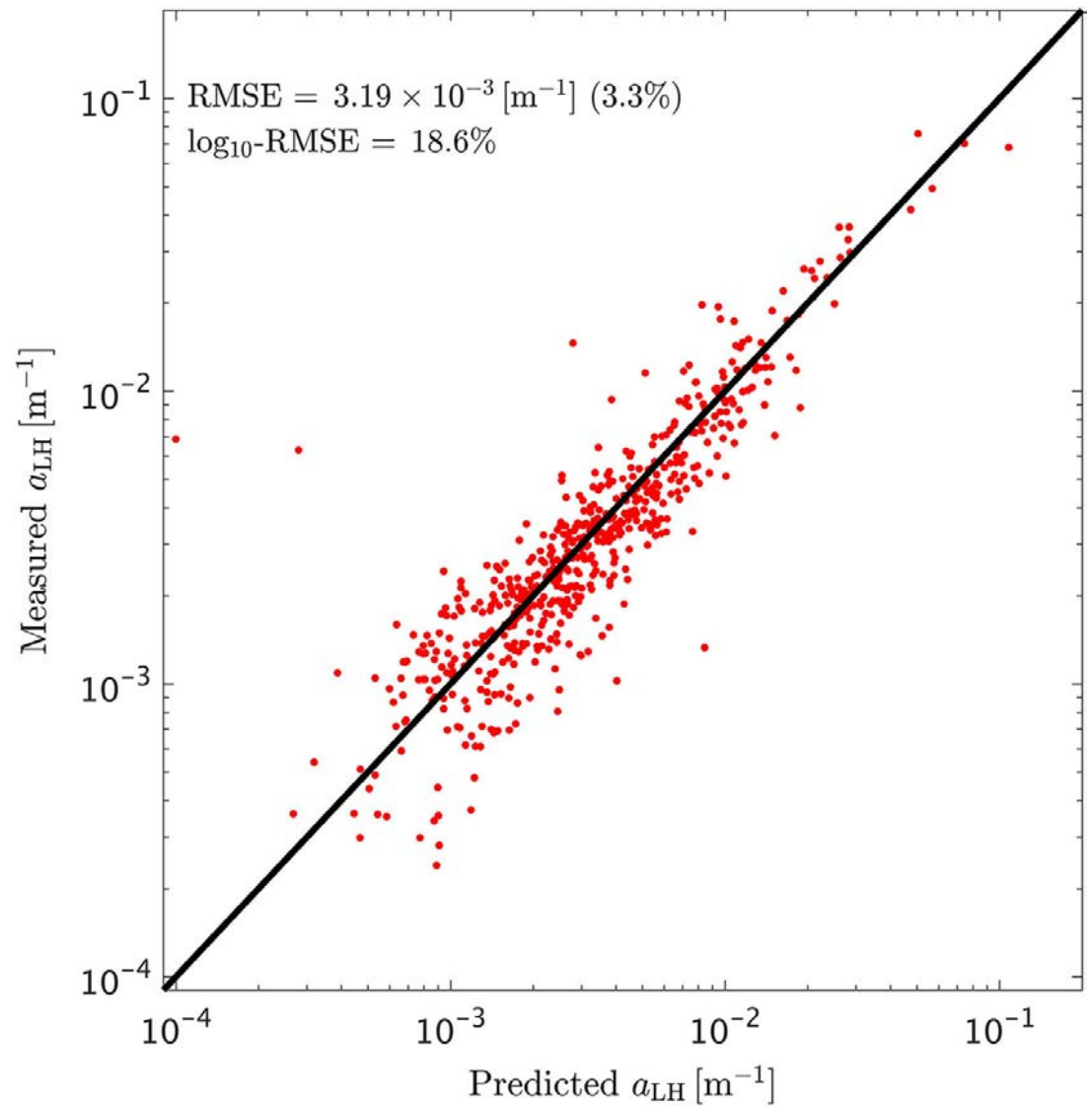


Roesler and Barnard, 2013

# Chlorophyll from absorption band effects on beam attenuation (Housekeeper et al., 2020)

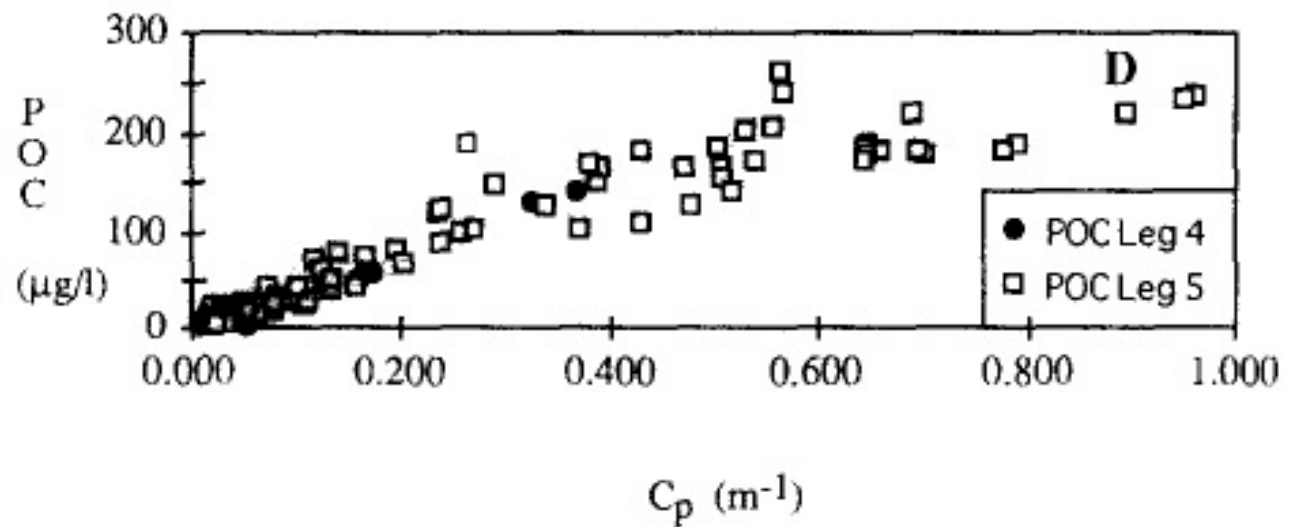


**Fig. 1.** (a) Anomalous diffraction approximation for  $Q_{\text{ext}}(\lambda)$  at a narrow absorption band as a function of light frequency for various  $\rho$  values, recreated from van de Hulst [8] using a lookup table; and (b)–(d) illustrative examples of  $Q_{\text{ext}}(\lambda)$  residuals for various small-sized phytoplankton (diameters 8, 5, and 1 μm, respectively), with a fixed  $n$  of 1.0344 and spectral  $n'$  with a maximum value of 0.0024 at the Chl *a* red absorption peak. The sizes presented in (b)–(d) are sensitive to the selection of real and imaginary refractive index; for example, as [14] illustrates, an anomalous dispersion curve for  $Q_b(\lambda)$  using a 1 μm absorbing sphere and spectral dependencies in both  $n$  and  $n'$ .



# Particulate Organic Carbon & Suspended Particulate Material proxy

- Backscattering and attenuation are associated with particle concentration / size.
- Backscattering is also highly dependent on morphology and type of the particle
- Carbon density in all oceanic particles / phytoplankton are not the same.



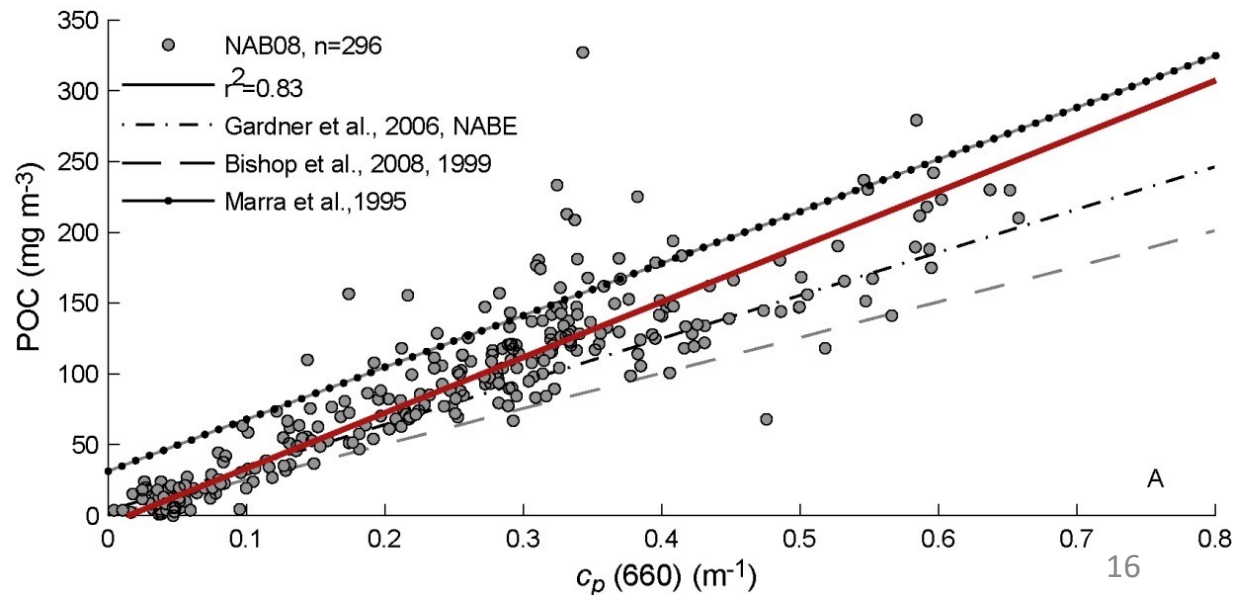
Gardner et al 1993

Also see:

Kitchen and Zaneveld, 1990

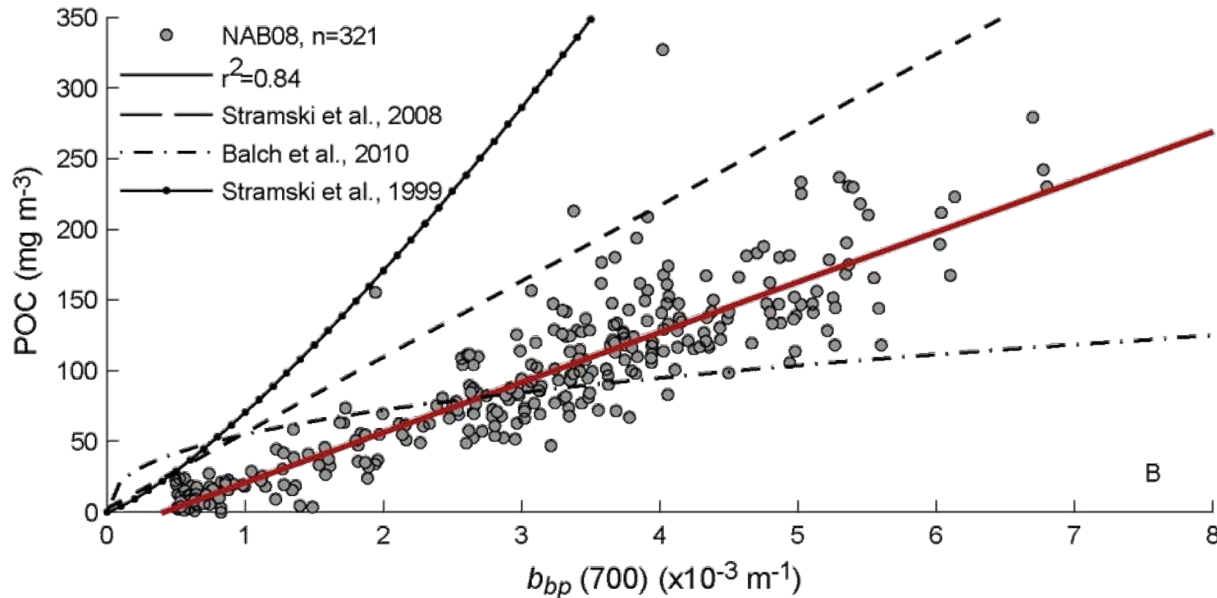
# Particulate Organic Carbon

Beam  
attenuation ( $c_p$ )



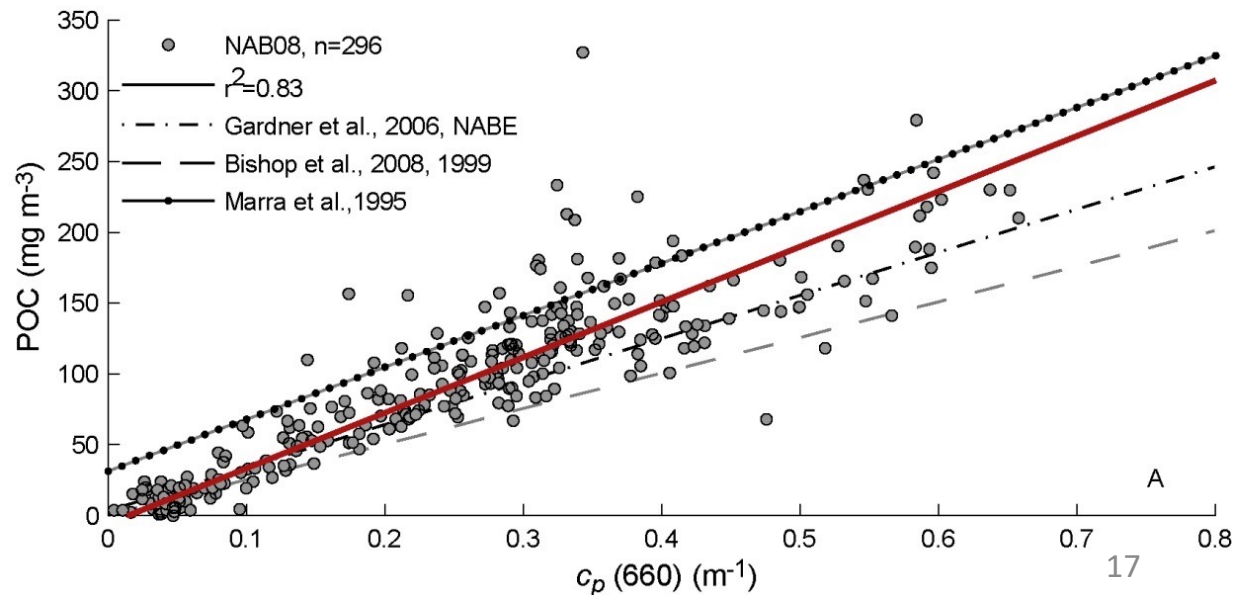


# Particulate Organic Carbon

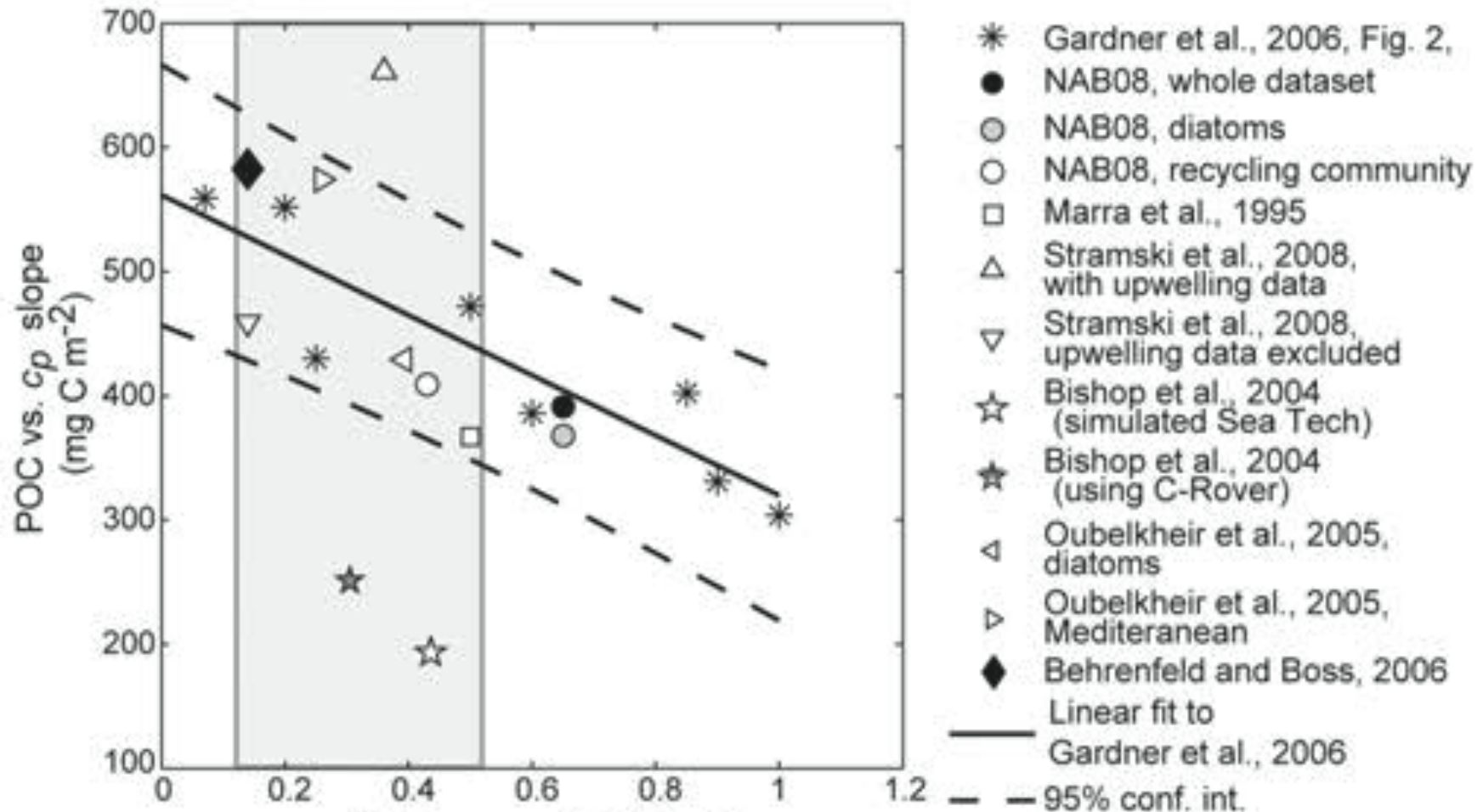


Particulate  
backscattering ( $b_{bp}$ )

Beam  
attenuation ( $c_p$ )



# POC/ $c_p$ slope comparison ( $\text{mg C m}^{-2}$ )

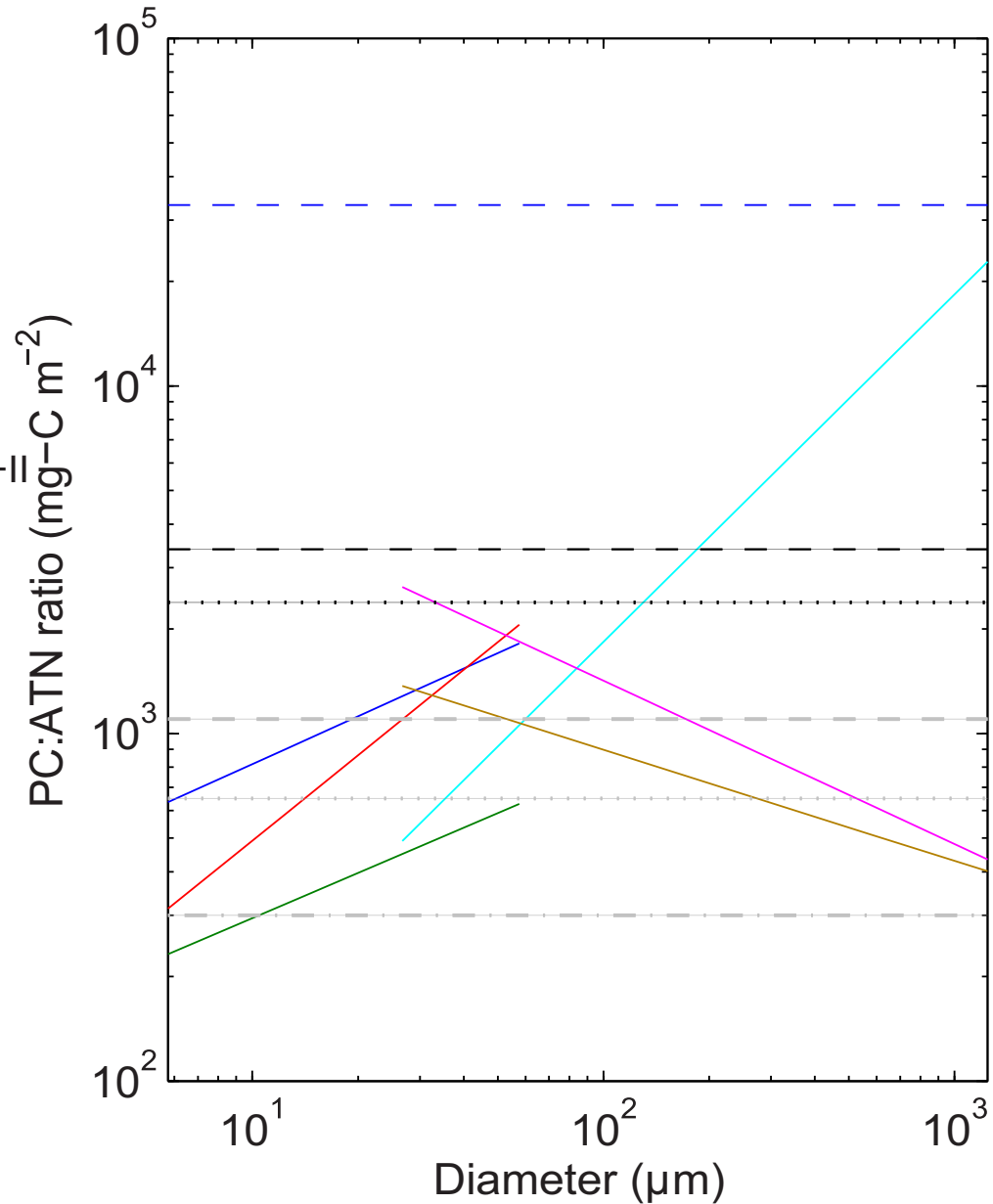
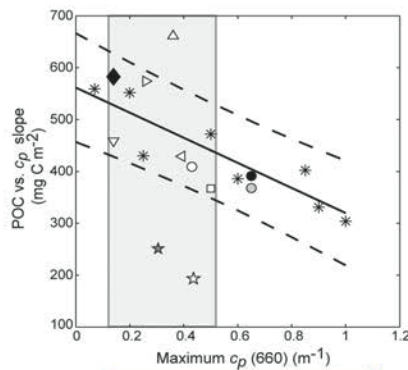


Increase in  $c_p$  & Chl  
 Smaller to larger cells  
 Oligotrophic to Eutrophic system  
 Pico/nano plankton to micro plankton

*Cetinić et al., 2012, JGR*

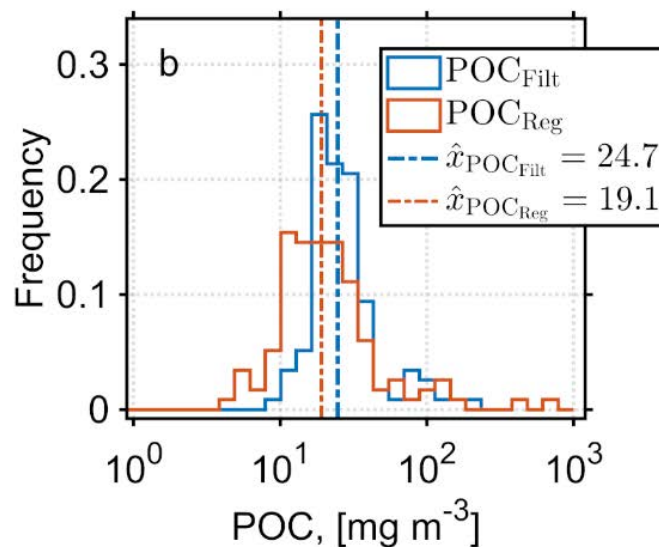
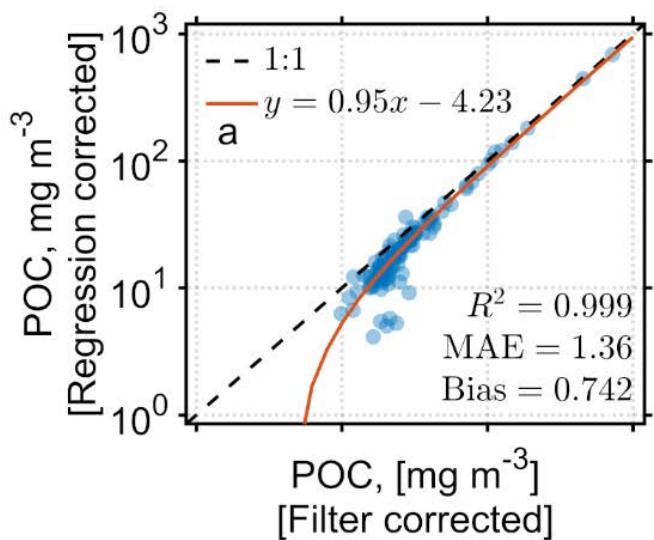
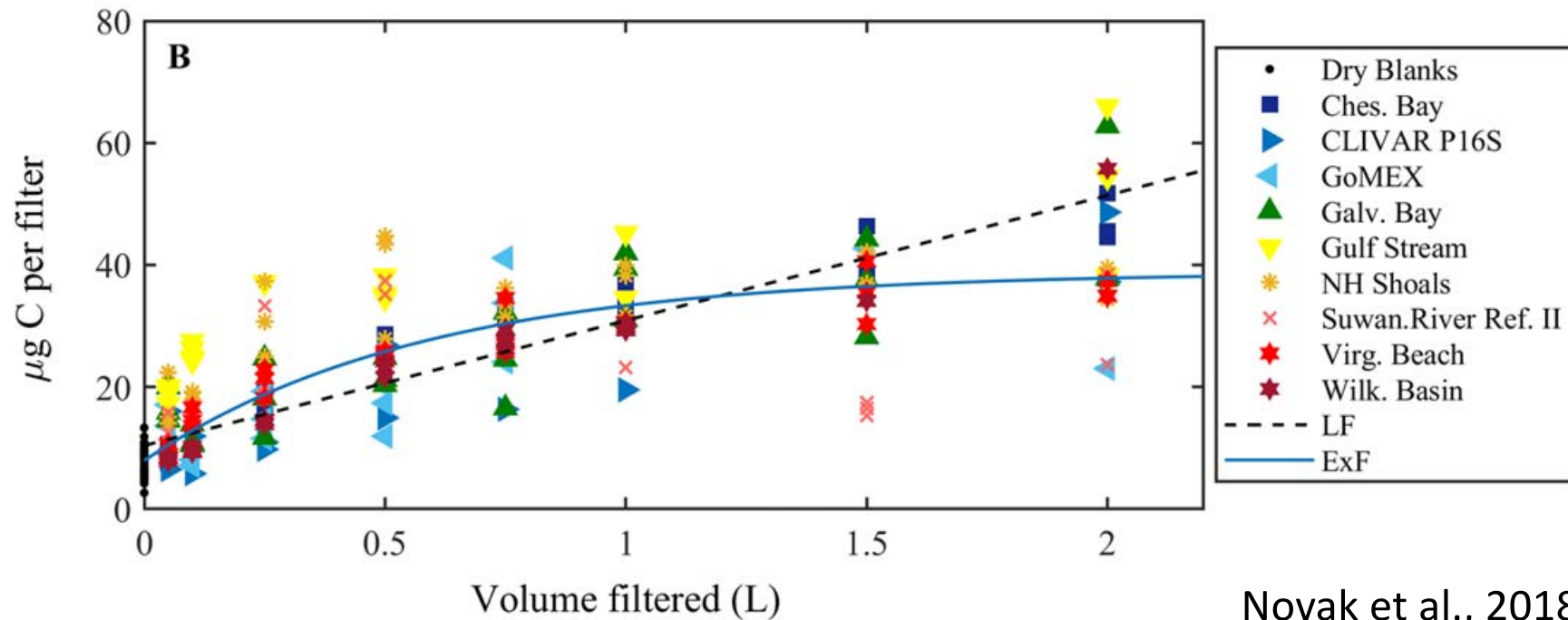
# Extend POC:cp relationship to sinking particles

- MD&L00, dinoflagellates
- MD&L00, diatoms
- MD&L00, protists
- D14 average fecal pellets
- A98 fecal marine snow
- A98 diatom marine snow
- - - B16 aggregates
- - - This study, weighted Type-II
- ⋯ This study, zero-intercept
- - - This study, T. weiss
- ⋯ This study, Iselin Pier
- - - This study, Saratoga L.



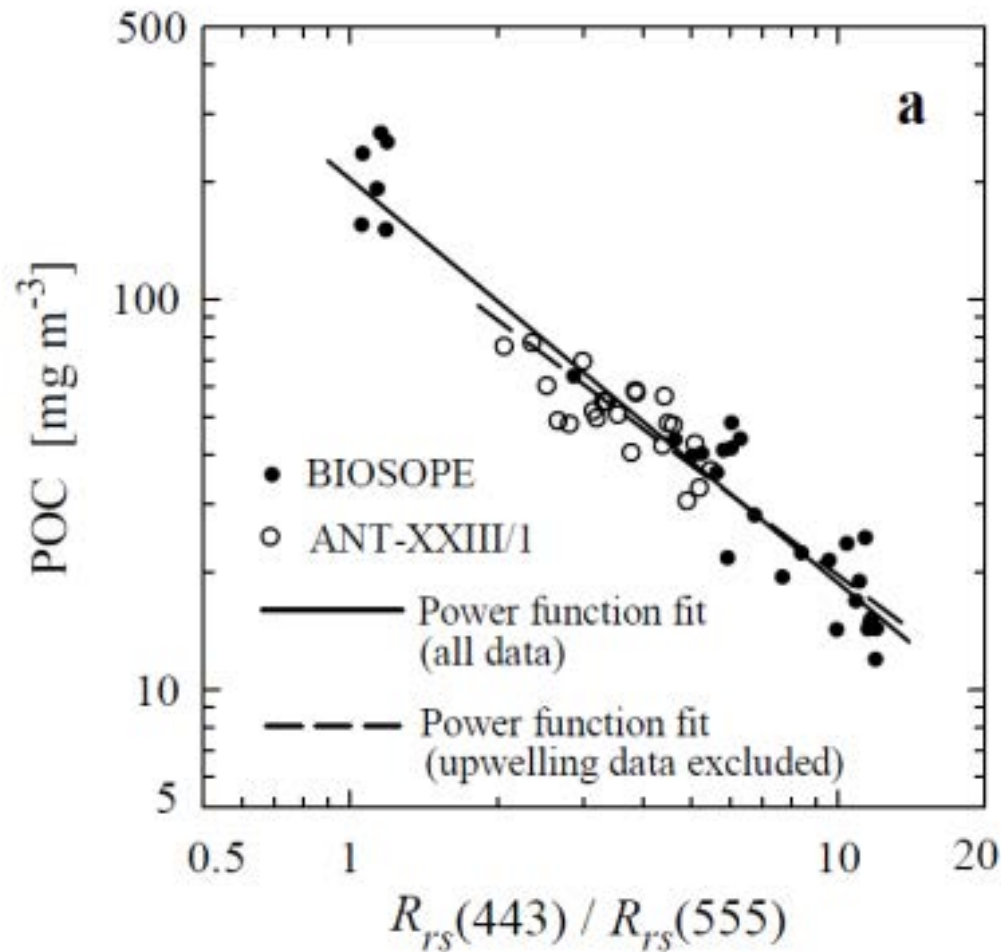
# Good POC proxies require accurate in-water measurements:

Bias in discrete POC measurements from DOC adsorption to filters



Chaves et al., 2021

# Current POC algorithm



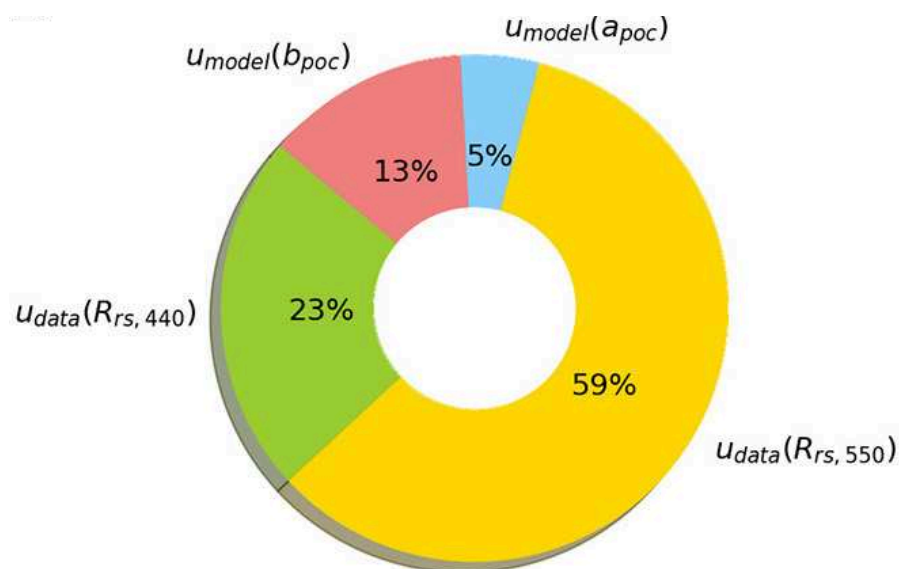
Stramski et al., 2008

No DOC adsorption blanks

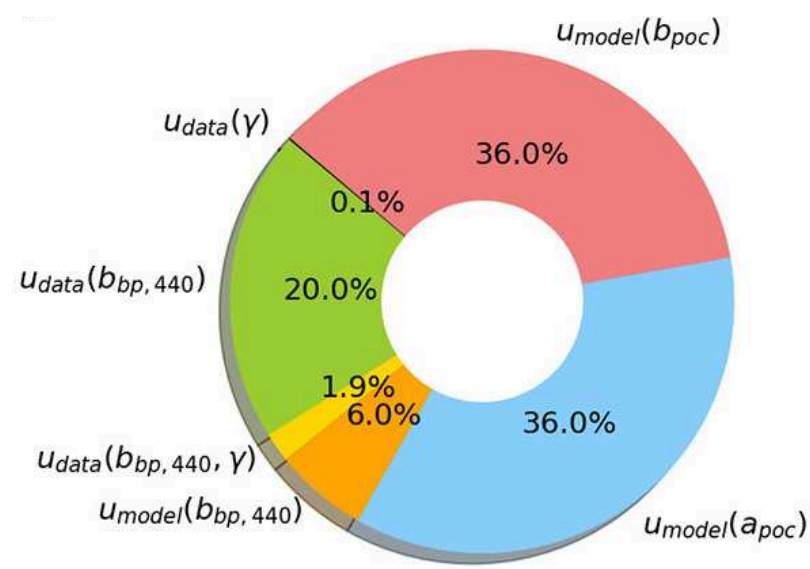
- DOC adsorption can be up to  $30.86 \pm 0.62 \mu\text{g C}$  per filter ([Novak et al. 2018](#))
- We need original data
- We need filtered volumes to apply the correction

# Contributions of different uncertainty sources to RS algorithms – example for two different POC algorithms

Stramski et al. 2008 band-ratio



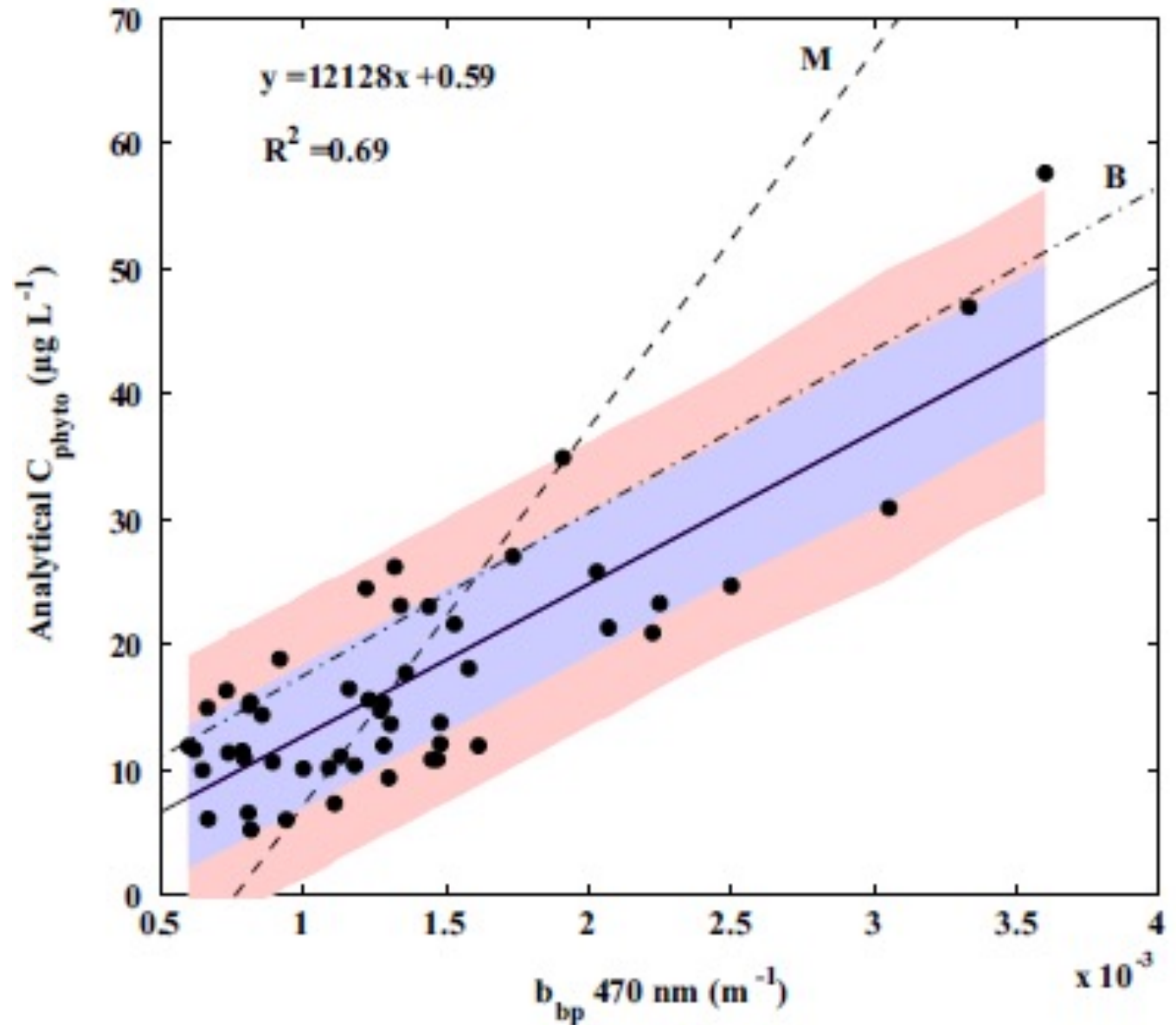
Rasse et al. 2017 IOP-based



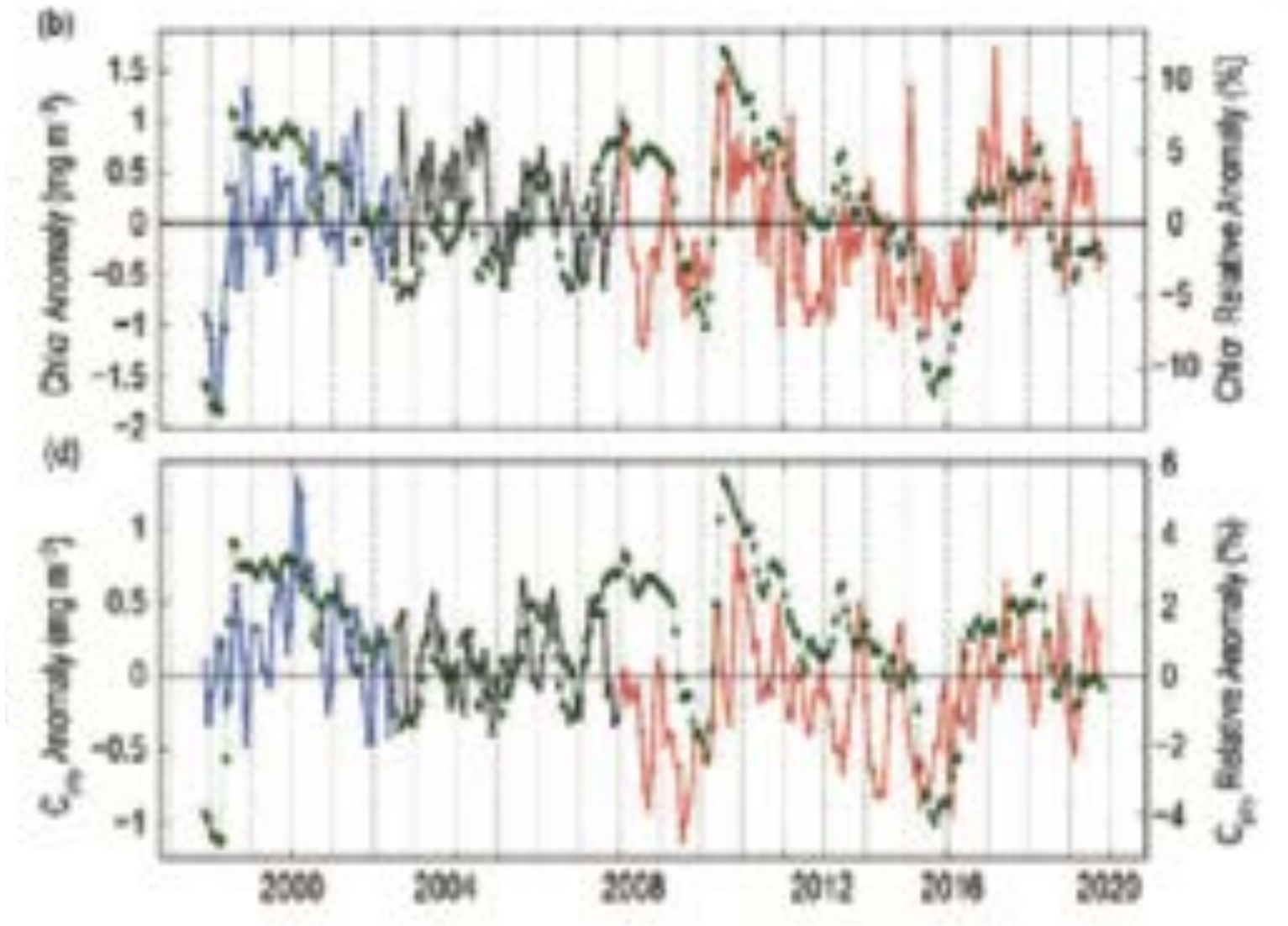
- $u_{model}$  terms include uncertainty in the POC measurements used to construct the algorithms (here, uncorrected DOC adsorption)
- $u_{data}$  terms include uncertainty in radiometric measurements

# Phytoplankton carbon

- Cell sorting technique in combination with optics
- Traditionally – calculation from of imaging/flow cytometry based biovolumes and cell/C values



# GIOP backscattering --> Phytoplankton carbon



Green dots = Multivariate ENSO Index, inverted

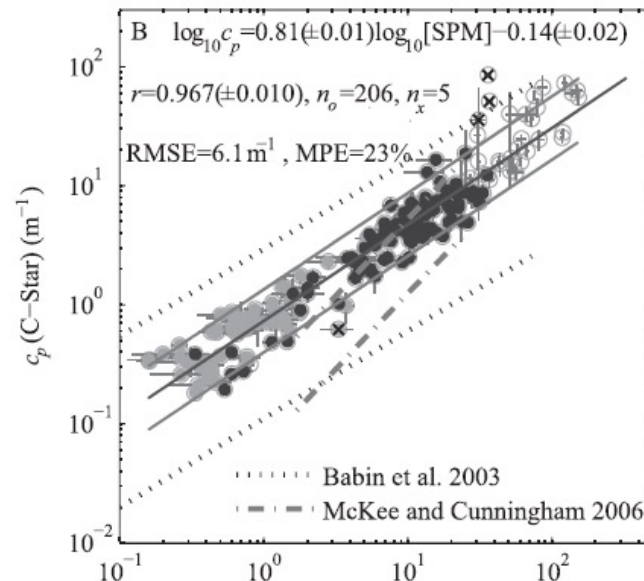
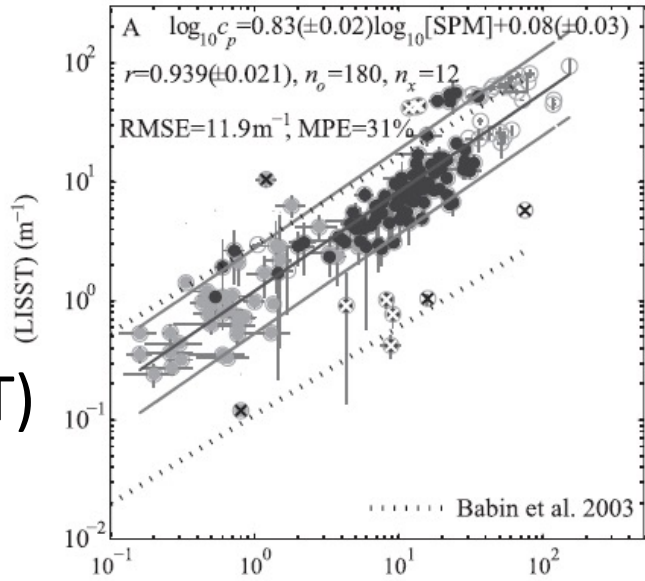
Franz et al. 2020, "Global Ocean Phytoplankton", in Baringer et al. [eds], *State of the Climate in 2019*



# Suspended Particulate Matter

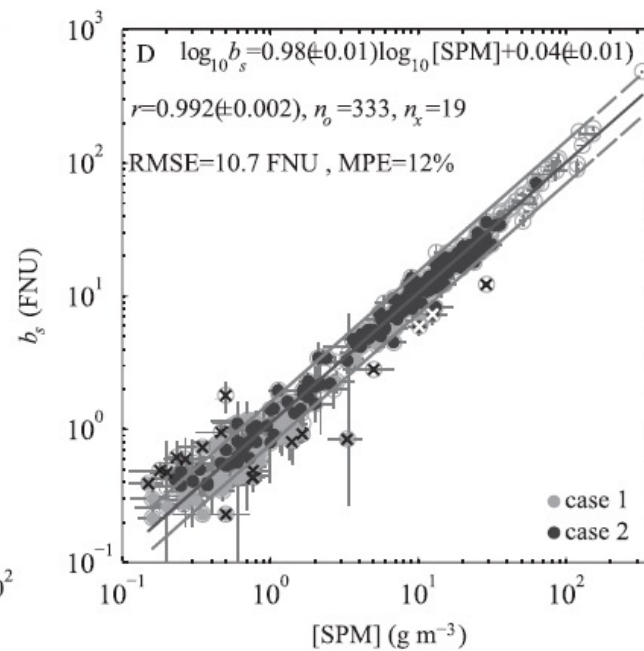
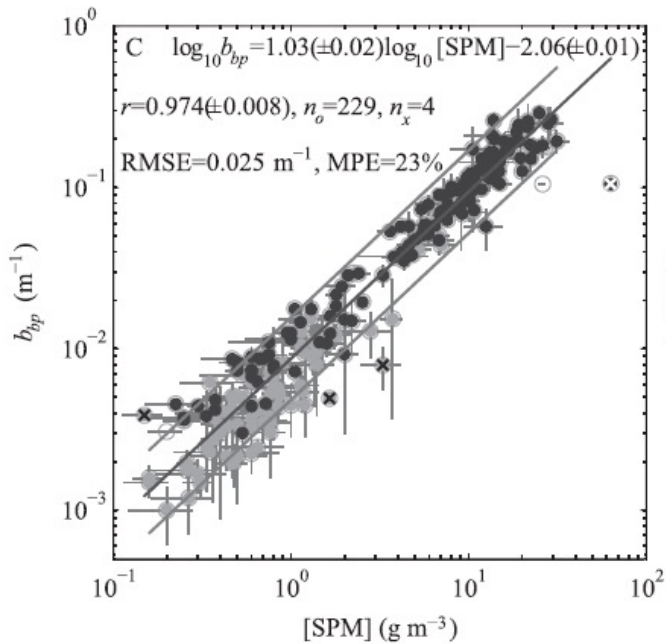
Gray dots: Case I  
Black dots: Case II

$c_p(\text{LISST})$



$c_p(\text{C-Star})$

$b_{bp}$



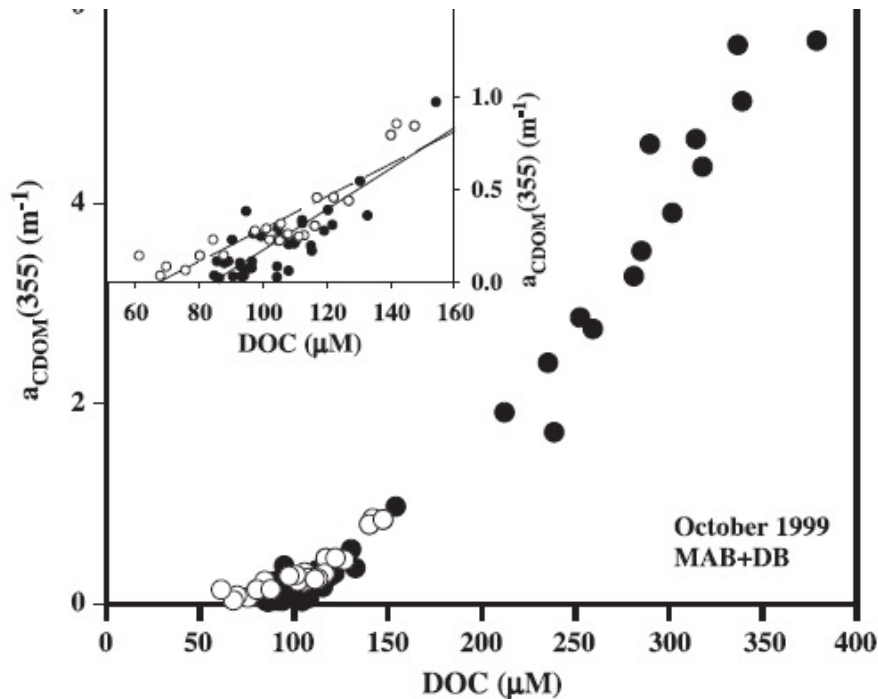
$b_s$  (discrete samples in turbidity sensor)

**SPM ( $\text{g m}^{-3}$ )**

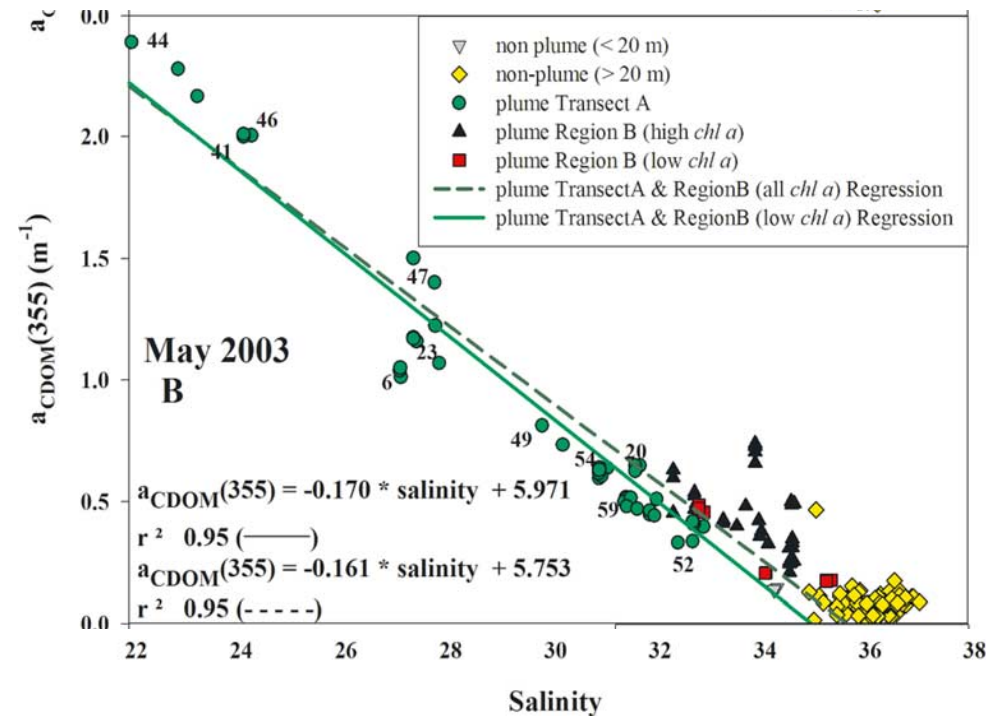
*Neukermans et al. 2012*

# Dissolved Organic Carbon

## CDOM can approximate DOC in the coastal ocean



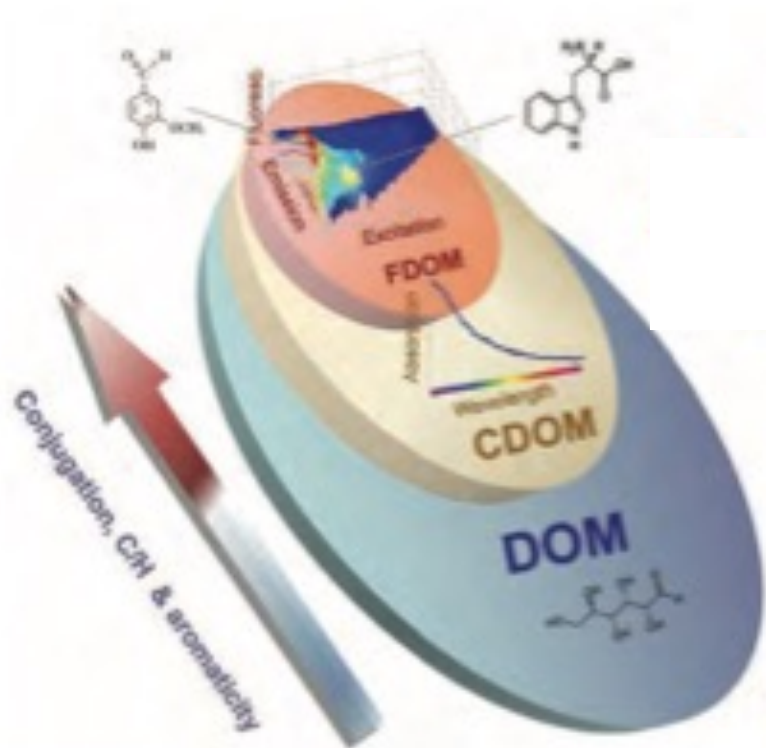
*Del Vecchio and Blough, 2004*



*Del Vecchio and Subramaniam, 2004*

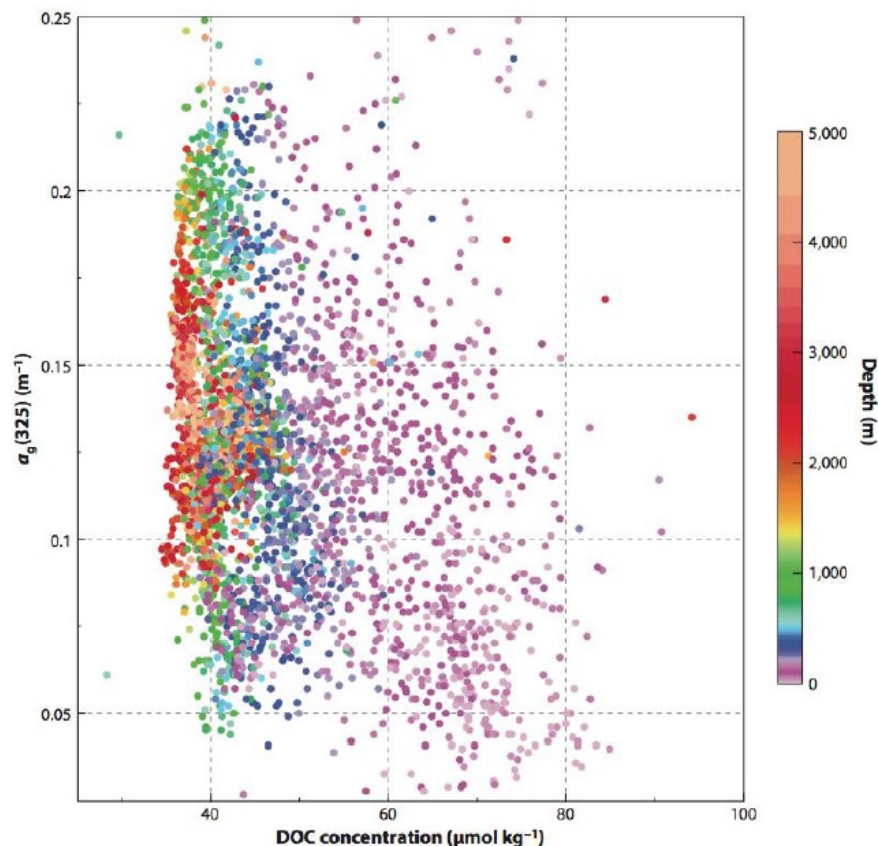
- Every coastal region will probably be different
- Better relationships when there is one strong source (e.g. river plume) and one major loss process (e.g. dilution into ocean)

# Dissolved Organic Carbon...



Stedmon and Alvarez-Salgado, 2011

## CDOM $\neq$ DOC in the open ocean



*Nelson and Siegel 2013*

DOC from space – [Aurin et al., 2018](#).  
Empirical regression against 4 wavebands (440-555 nm) + salinity  
required, RMSE still 27-29  $\mu\text{mol L}^{-1}$

# Overview

- ✓ Where proxies fit, why to construct them
- ✓ Concentration proxies
- Composition and rate proxies
  - “ANYTHING THAT CAUSES VARIABILITY IN THE SAMPLE IS AN OPPORTUNITY TO EXTRACT ADDITIONAL INFORMATION FROM THAT SAMPLE” - COLLIN ROESLER
- Caveats

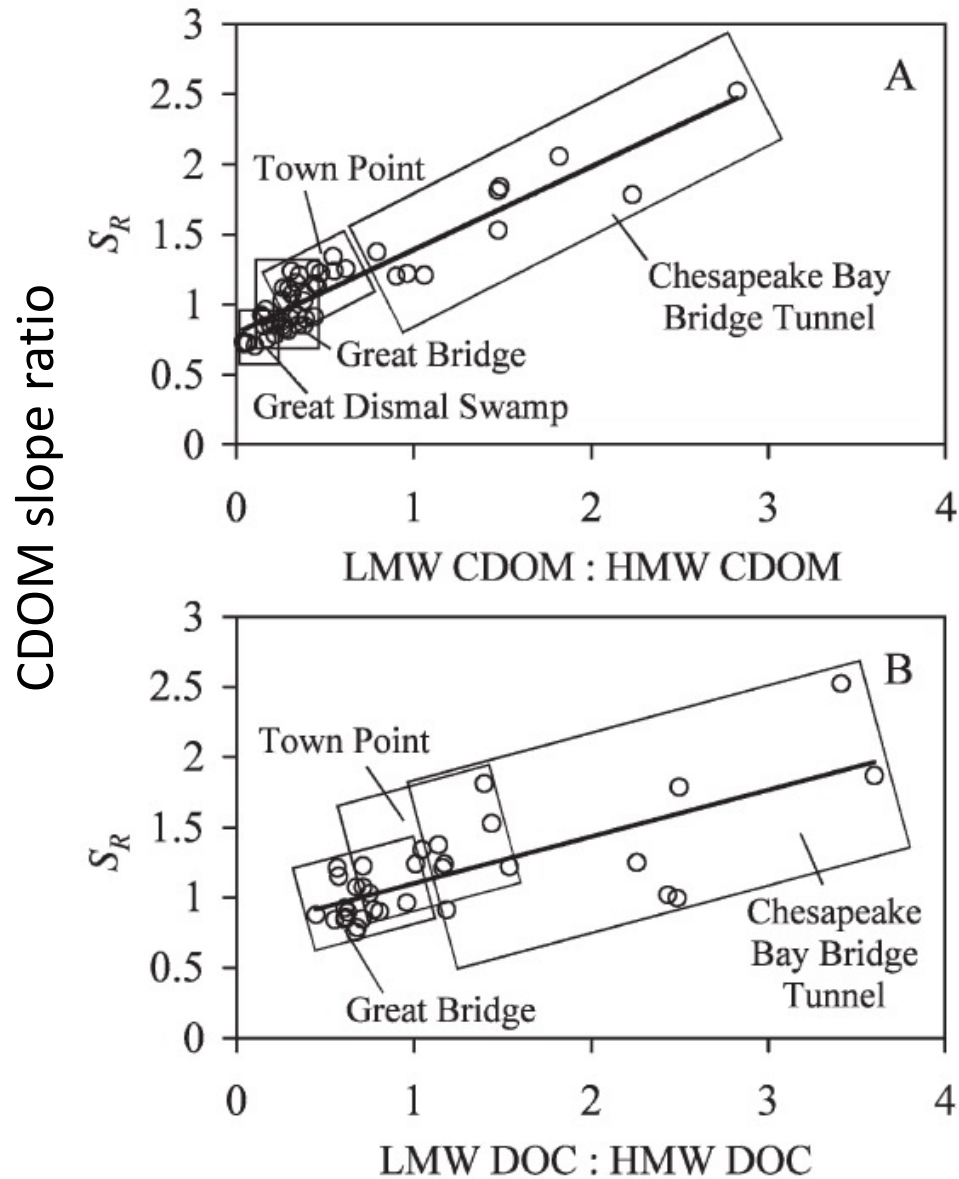


As Rufus well knows, there's opportunity  
in chaos.

# OPPORTUNITY IN CHAOS

## QUALITY (COMPOSITION, SIZE)

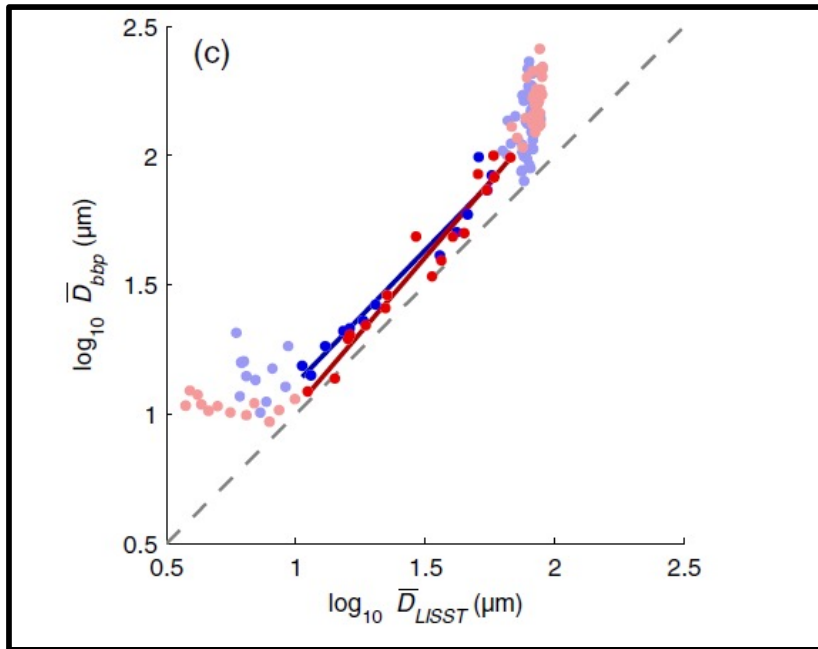
# CDOM slope $\sim$ DOC molecular mass



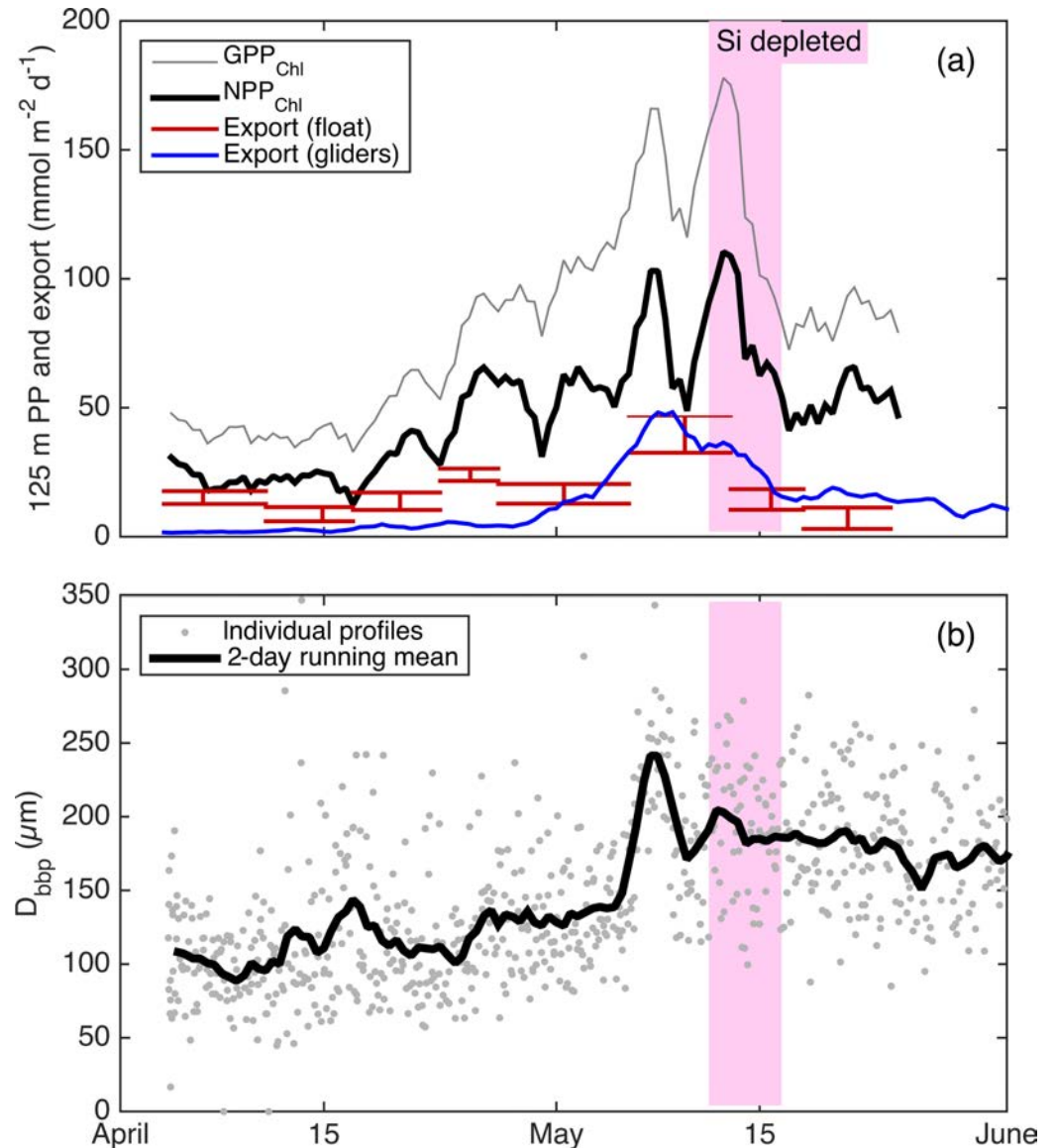
[Helms et al., 2008](#)

# Particle size – fluctuation based

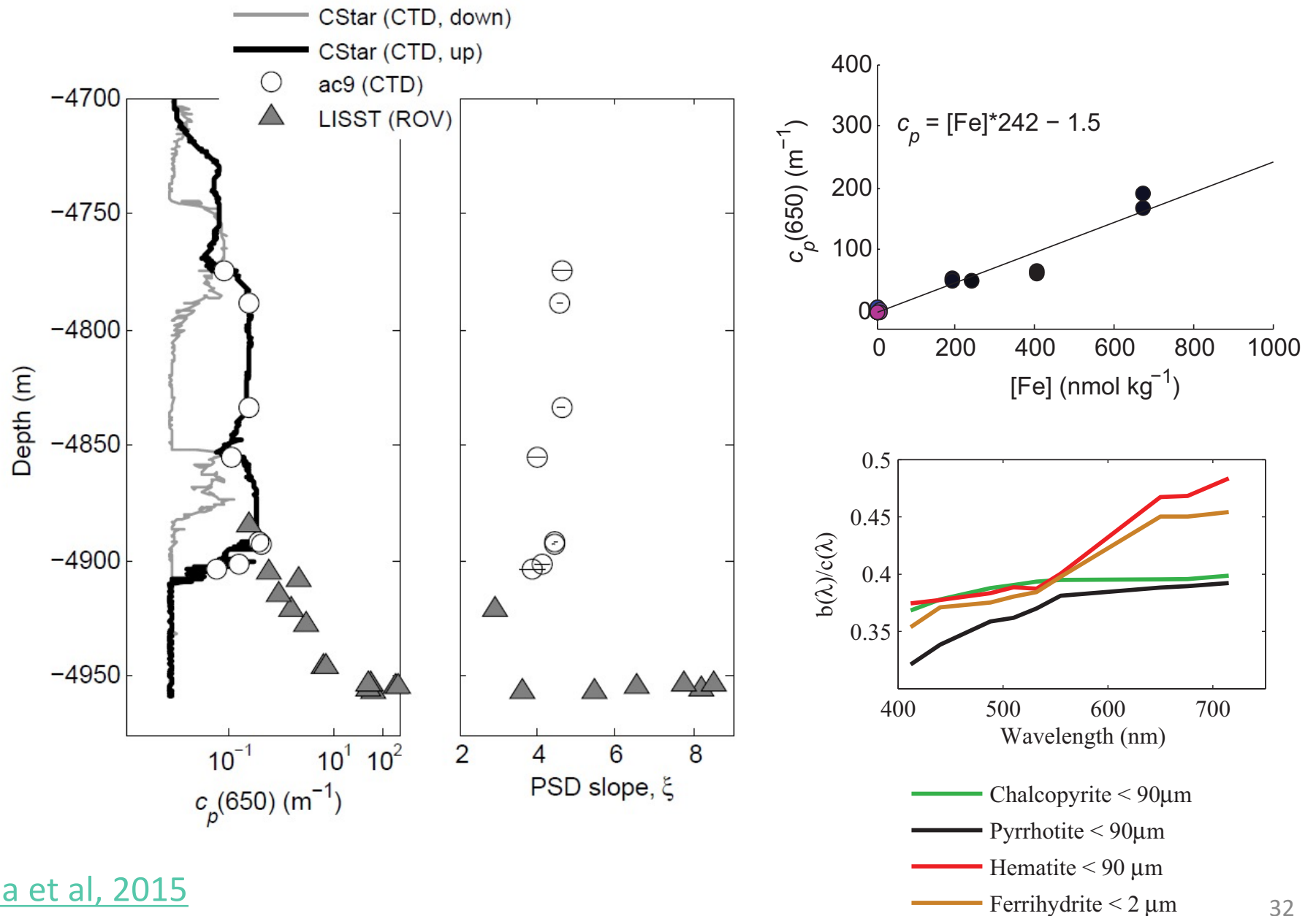
Application – Briggs et al., 2018



Briggs et al, 2013

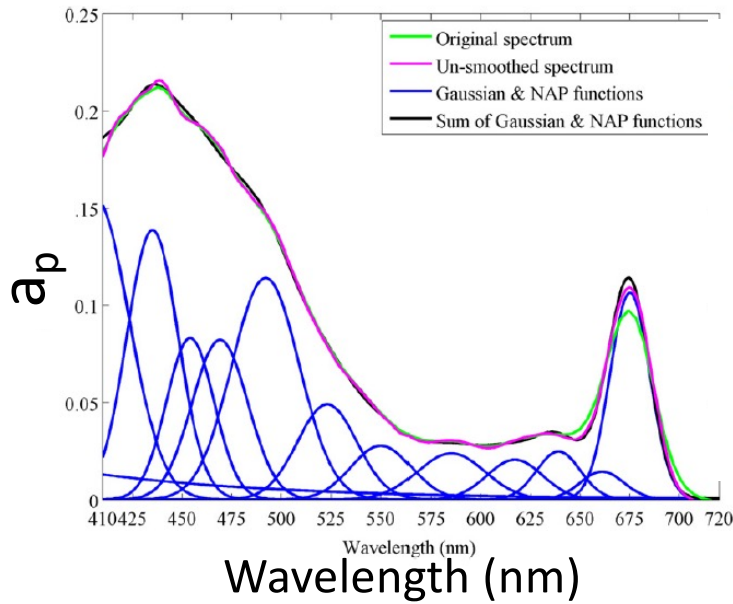


# Particle composition in a hydrothermal plume





# Phytoplankton community composition

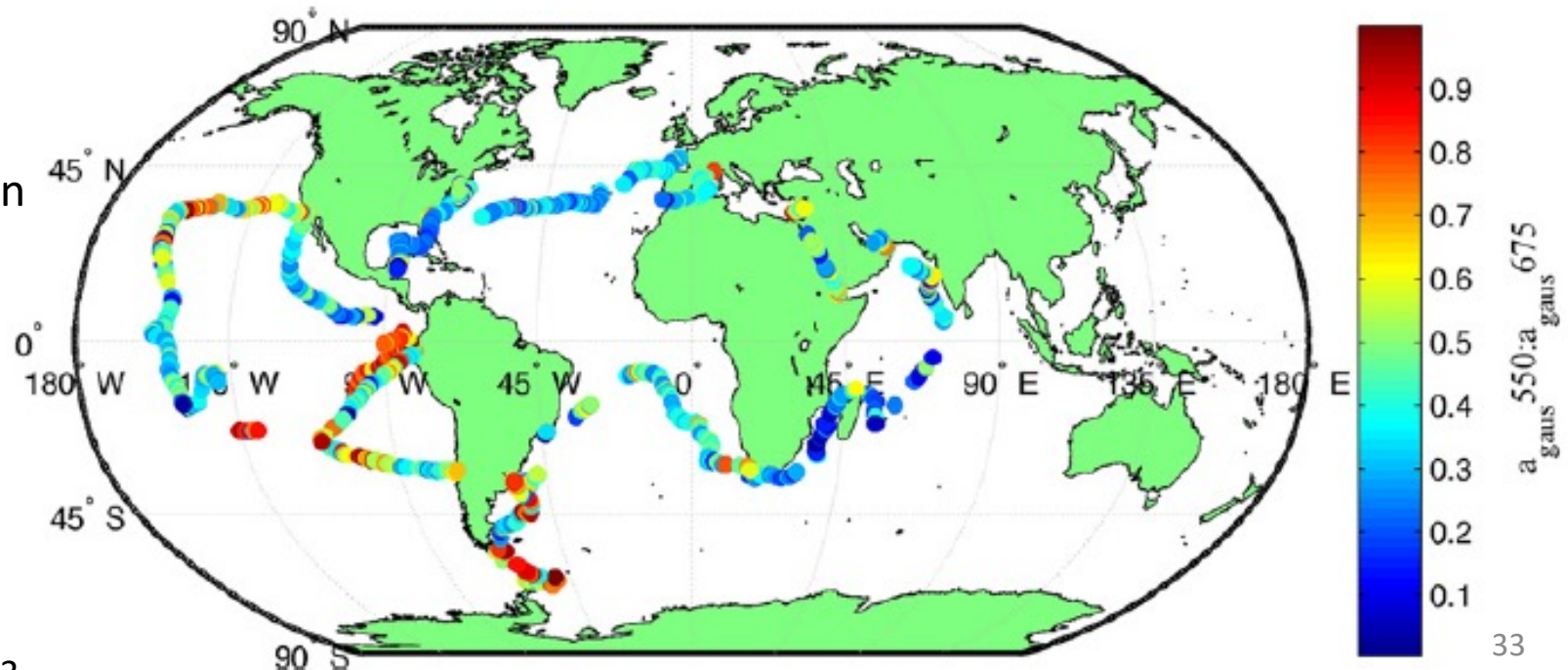


**Table 2**

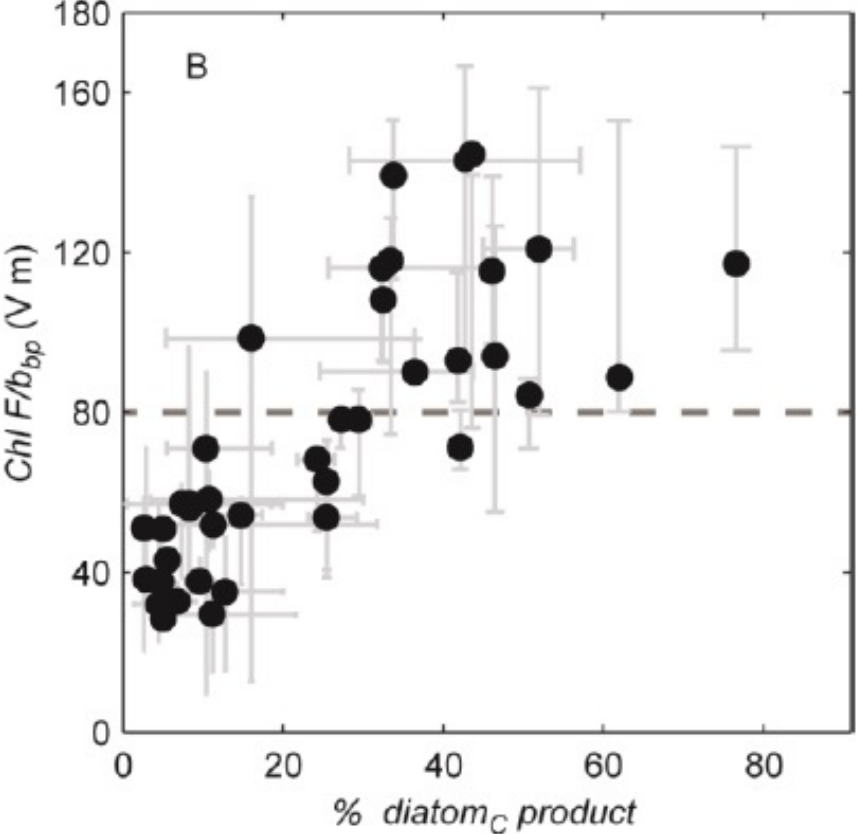
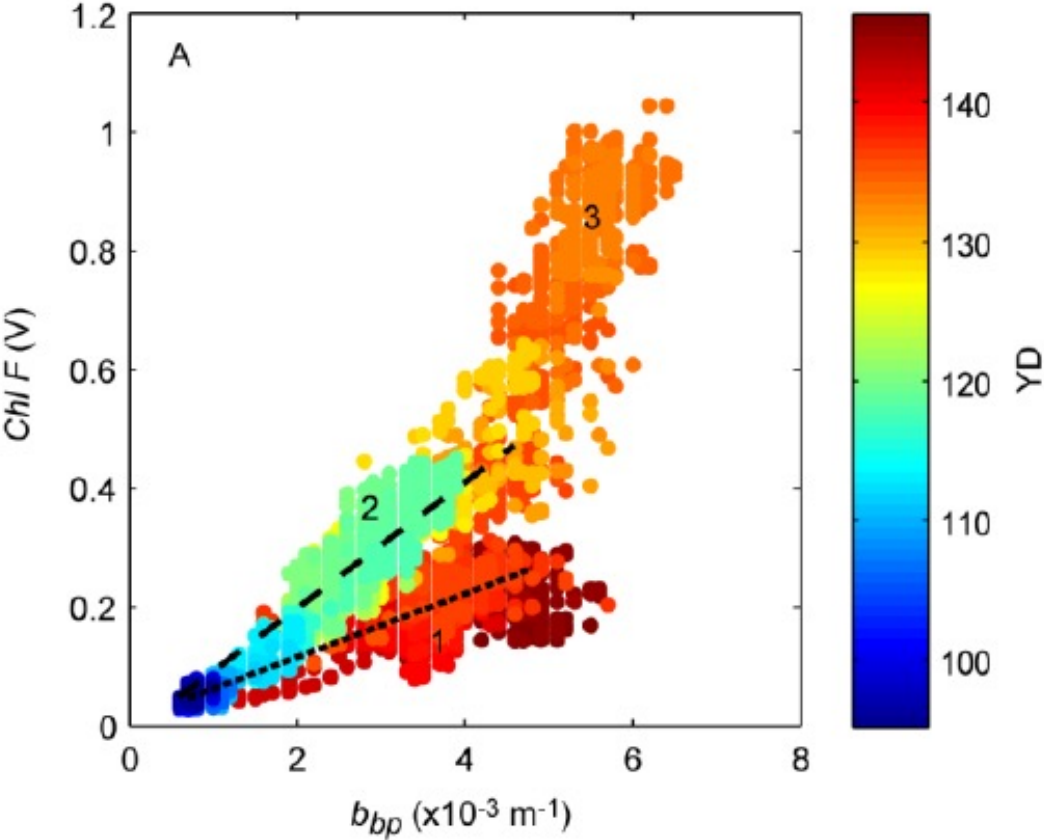
Correlations between HPLC pigment concentrations and  $a_{\text{gaus}}(\lambda_i)$  at ten different pigment absorption wavelengths. Correlation values are Spearman's rank correlation coefficient (non-parametric; denoted  $\rho$ ).  $A$  and  $B$  are coefficients determined using Eq. (4) (Section 2.4).

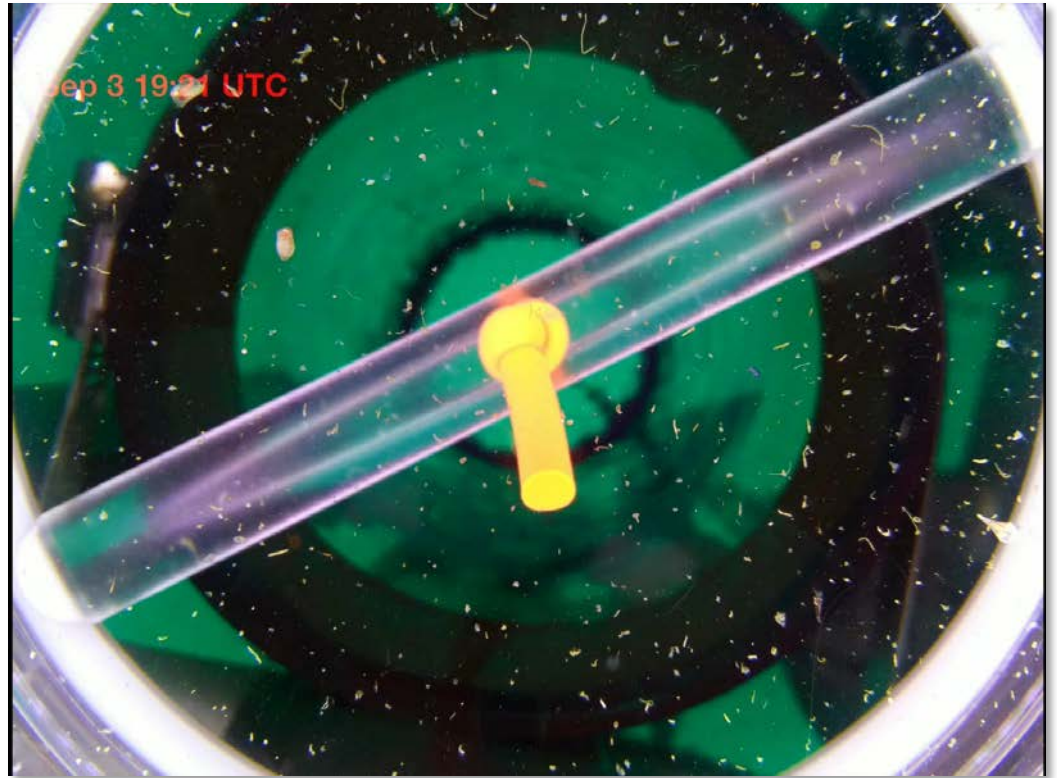
Wavelength (nm)	Pigment(s)	$\rho$	$A$	$B$	$e_{\text{median}} (\%)$
435	TChl <i>a</i>	0.868	0.031	0.578	35
617	TChl <i>a</i>	0.834	0.003	0.758	36
675	TChl <i>a</i>	0.899	0.014	0.798	30
454	0.03(TChl <i>b</i> ) + 0.07(Chl <i>c</i> )	0.845	0.028	0.414	57
469	TChl <i>b</i>	0.783	0.066	0.533	52
661	TChl <i>b</i>	0.747	0.018	0.668	40
585	Chl <i>c</i>	0.846	0.014	0.582	53
639	Chl <i>c</i>	0.894	0.012	0.641	41
492	PPC	0.606	0.046	0.650	51
523	PSC	0.855	0.013	0.588	49

Phycoerythrin proxy distribution



# Phytoplankton Community composition





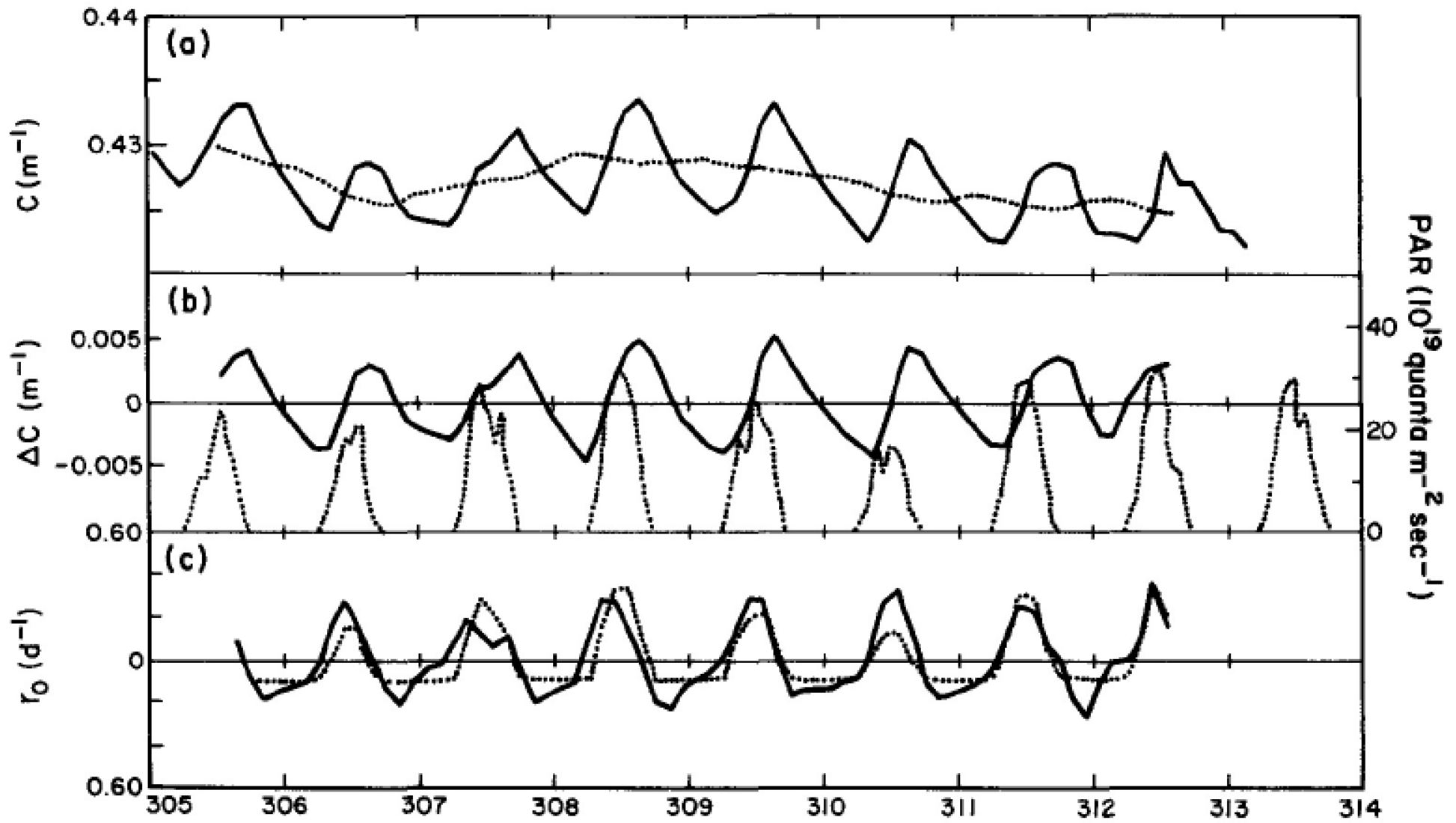
Omand et al, unpublished

# OPPORTUNITY IN CHAOS RATES AND FLUXES

Change in *in situ* optical proxies will tell us something about **rates and fluxes**

- Space - e.g. carbon export from mixed layer to deeper ocean
- Time – productivity - e.g. primary production
- Type – e.g. phytoplankton to detritus, POC to DOC
- Typical units could be:
  - $\text{mol m}^{-2} \text{y}^{-1}$
  - $\mu\text{mol L}^{-1} \text{h}^{-1}$
  - $\text{d}^{-1}$

# PRODUCTIVITY/GROWTH



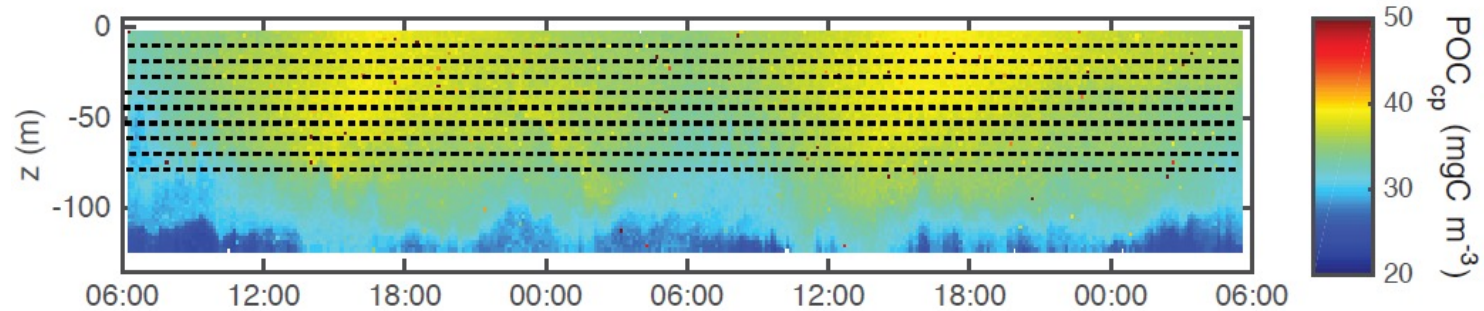
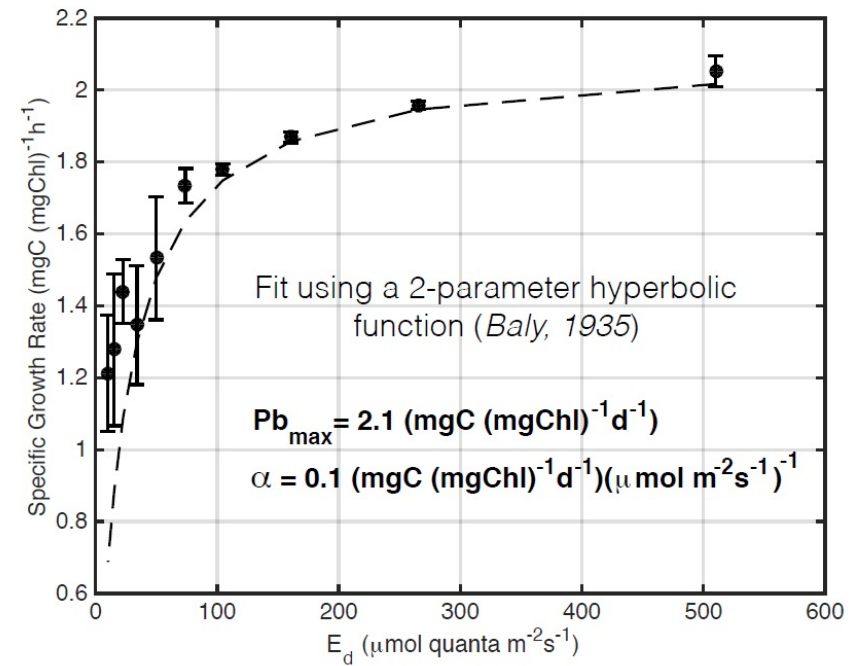
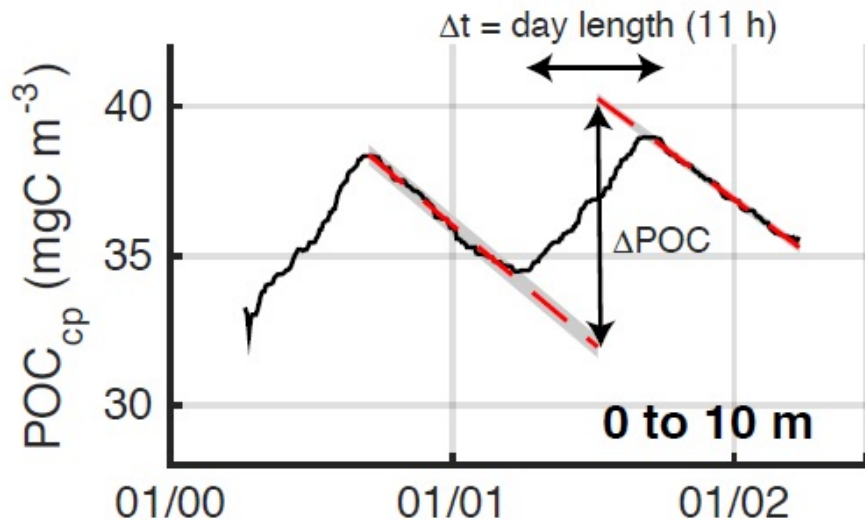
(at 20 m, from CTD profiles 8x/day)

LOCAL DAY 1982

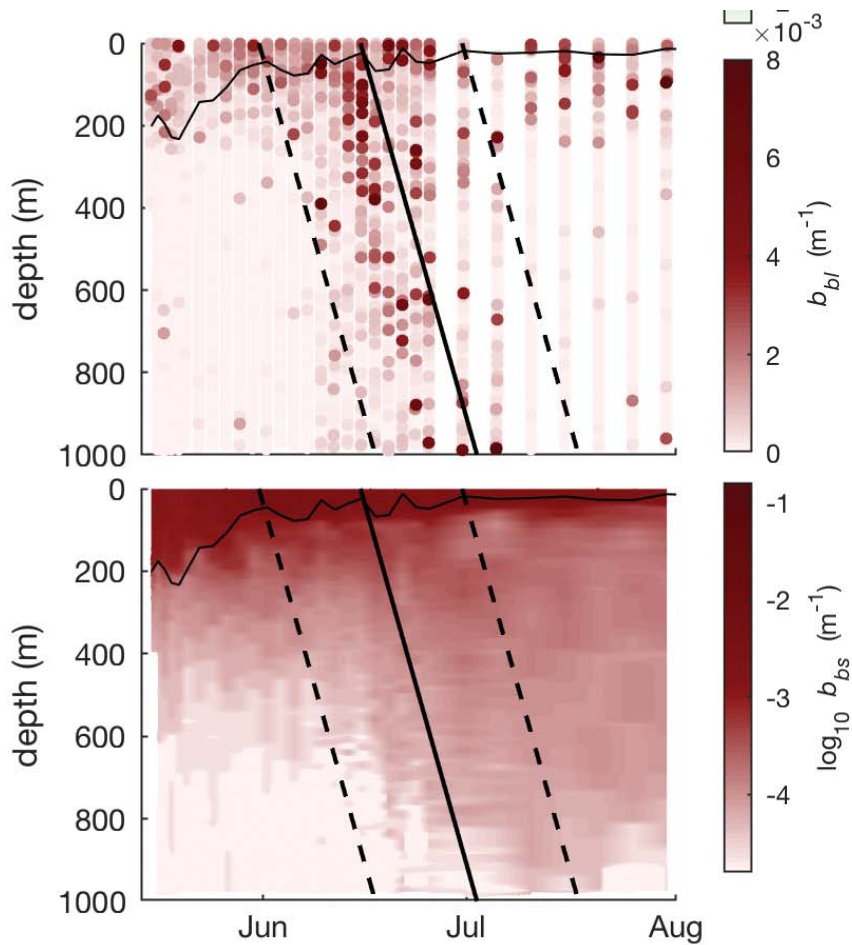
Siegel et al., 1989

+relevant work by many others

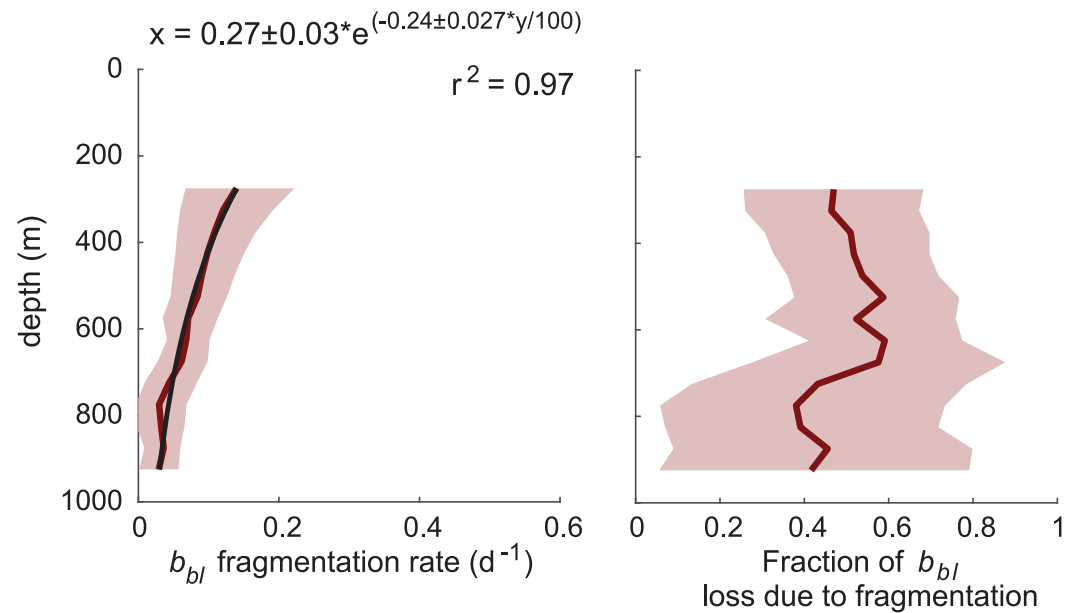
# PRODUCTIVITY/GROWTH



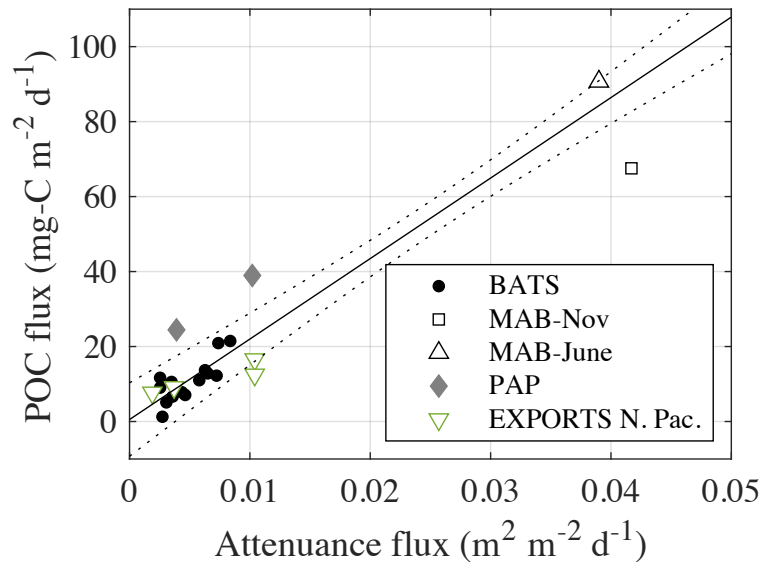
# DISSAGGREGATION



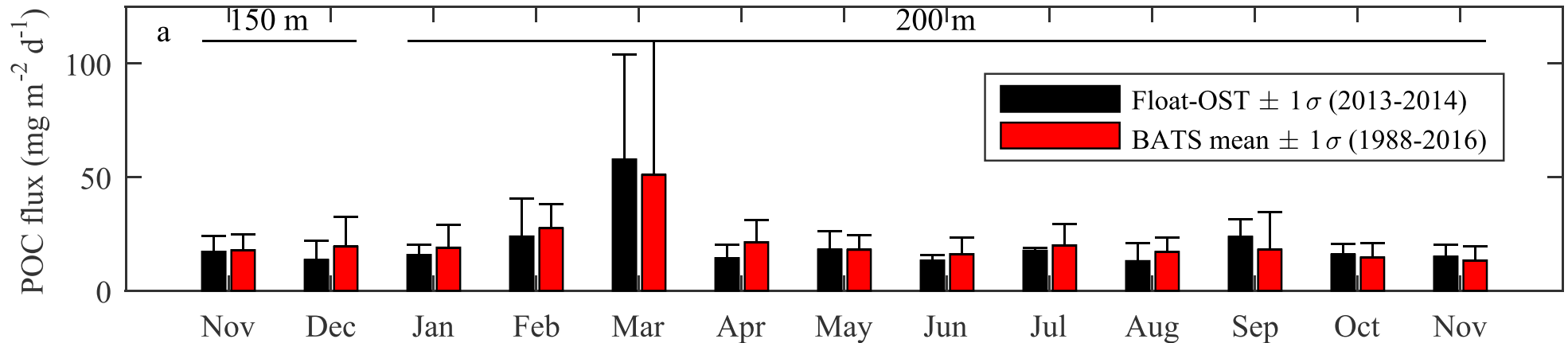
- Use “spikes” to quantify large- and small-particle backscattering
- Find well-defined export events in BGC-Argo data
- → Disaggregation during such events accounts for  $\sim$  half of C flux attenuation



# “Optical sediment trap”

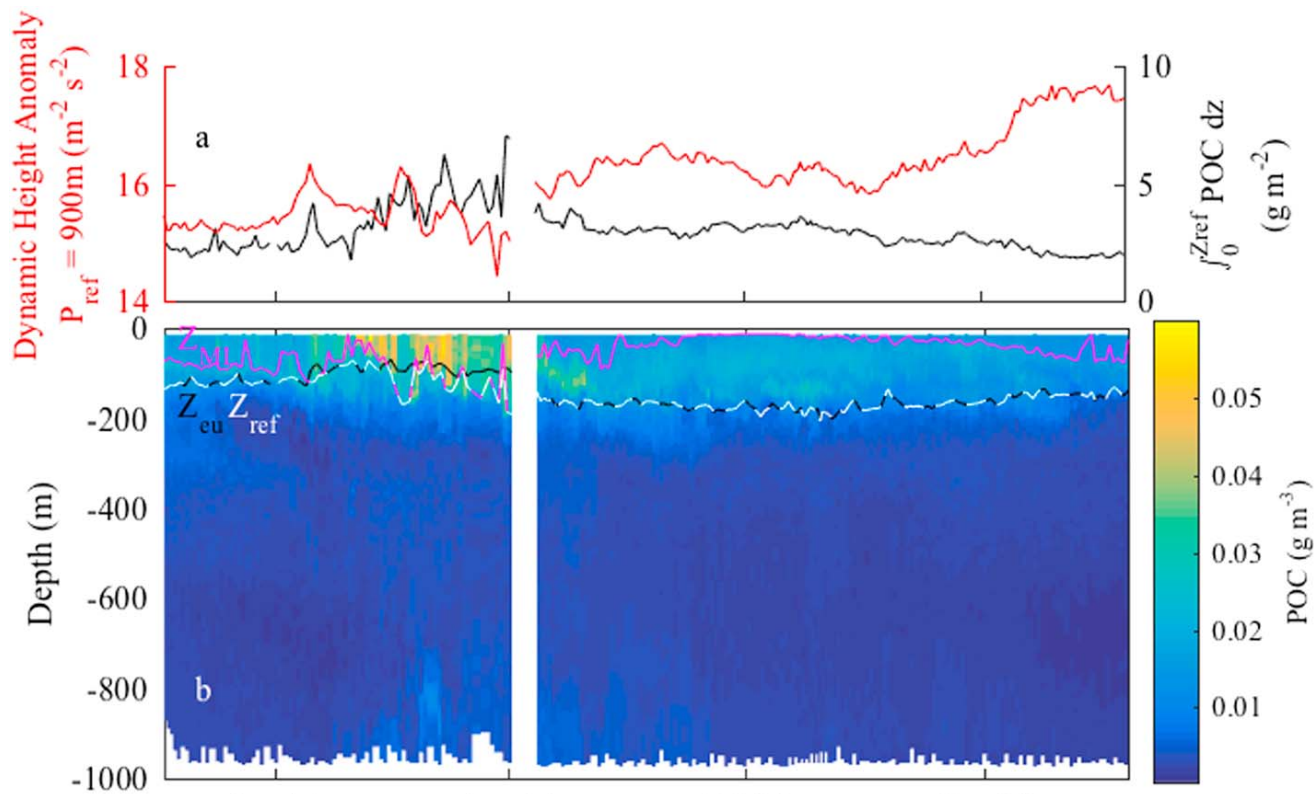


Estapa et al., 2017

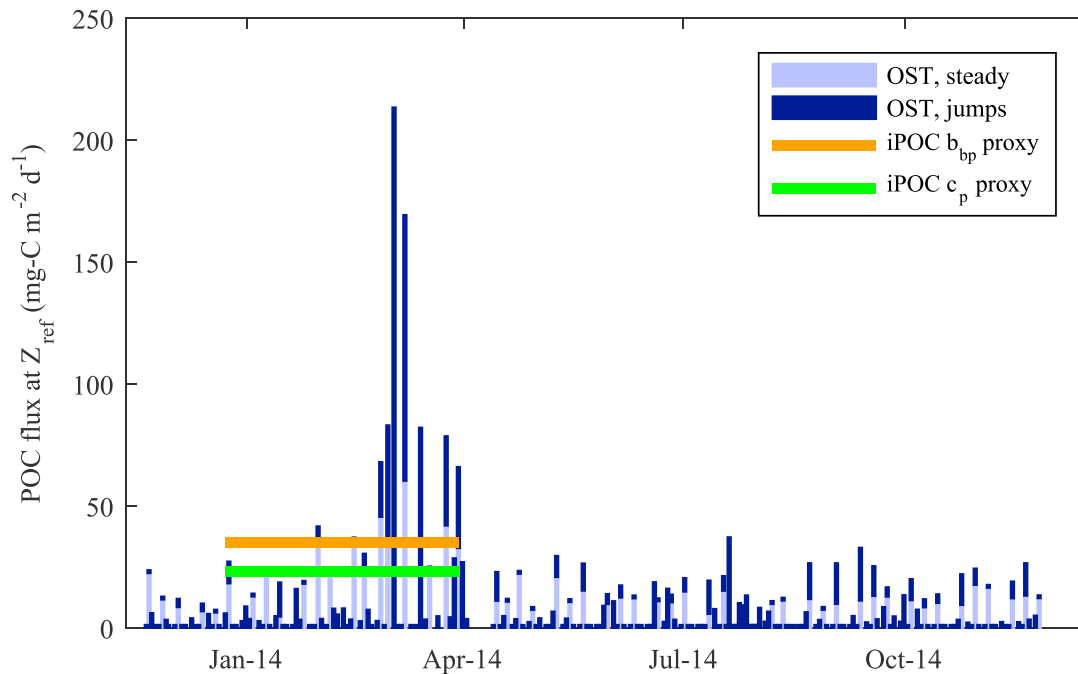


Estapa et al. 2019, plus unpublished EXPORTS data





“Optical  
 sediment  
 trap”



Intercomparison of three  
 optical proxies for POC flux

# Overview

- ✓ Where proxies fit, why to construct them
- ✓ Concentration proxies
- ✓ Composition and rate proxies
- Caveats

# Curt's slide from Monday morning...

## Take-home Messages

Statistical methods for retrieving environmental information from remotely sensed data have been highly successful and are widely used, but...

- An empirical algorithm is only as good as the underlying data used to determine its parameters.
- This often ties the algorithm to a specific time and place. An algorithm tuned with data from the North Atlantic probably won't work well in Antarctic waters because of differences in the phytoplankton, and an algorithm that works for the Yellow Sea in summer may not work there in winter.
- The statistical nature of the algorithms often obscures the underlying biology or physics.

# Caveats

1. Validate – make sure your proxies are based on strong and meaningful relationship with biogeochemical parameters
2. Interpolate rather than extrapolate – know the limits of your method, spatial, temporal and logical
3. Same as Rufus the dog, seize the variability and chaos (but remember 1 and 2)

“PROXIES WORK UNTIL THEY DON’T” –  
Mary Jane Perry



As Rufus well knows, there's opportunity  
in chaos.