Ocean Optics Summer Class

Calibration and Validation for **Ocean Color Remote Sensing**

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What is Closure?

Closure refers to obtaining the same quantity in at least two different ways.

If the agreement is good (within the uncertainties of the methods used) we say that closure has been achieved.

Performing closure in papers is a great way to convince reviewers to accept your paper—i.e., that you have good data and good modeling.

Data-data closure shows agreement between disparate data sets, e.g. using two different instruments and/or different calibration methodologies, AOP and IOP, in-situ vs. remote.

Model-model closure indicates that model assumptions (e.g. plane-parallel vs 3D; with and without inelastic scattering) do not introduce significant error.

Scale closure: are measurements made on small spatial scales (cm³ water samples; a mooring) consistent with large-scale measurements (*K*_d, *R*_{rs}; a satellite)

Model-data closure ties it all together

• Show two examples of output from advanced HydroLight simulations to show what is necessary to achieve modeldata closure, i.e., getting all of your inputs to H and outputs from H to agree with your measurements

Measurements Necessary for Model-Data Closure

HydroLight inputs

• absorption coef $a(z,\lambda)$ (e.g., from ac-9 or spectrophotometer) • scattering coef *b*(*z,*λ) (e.g., from ac-9)

• scattering phase function $\beta(z,\lambda,\psi)$ (almost never measured, but may have backscatter fraction $B=b_{\rm b}/b$ from $b_{\rm b}$ (e.g., HydroScat or EcoVSF) and *b* (ac-9)

• boundary conditions: sea state (wind speed); sun location and sky conditions (usually model), bottom reflectance (in shallow water)

HydroLight outputs

• radiometric variables (radiances and irradiances; usually measure *L^u* (*z,*λ) and *E^d* (*z,*λ) at a minimum) \cdot apparent optical properties (K_{d} , $\mathsf{R},$ $\mathsf{R}_{r\mathrm{s}}$ etc obtained from radiometric measurements). The most common for remote sensing is remote sensing reflectance R_{rs} (often measure E_{d} (air) and L_{u} (z) and extrapolate upward from underwater *L^u* , or estimate *Rrs* using above-surface techniques)

Comprehensive Data Sets Are Extremely Scarce

Everyone wants comprehensive data sets, but no one wants to pay for them, and scientists don't want to be forced to take data they themselves aren't going to use.

See my notes on comprehensive data sets in the Papers directory (ComprehensiveDataSets.pdf) for an overview of what should be measured in a field experiment, but never is (cost, lack of interest, ignorance, politics, …)

When you go home and design your grand field experiment, at least look at these notes and do the best you can with the available resources.

The HyCODE Data Set

Data set from ONR HyCODE (Hyperspectral Coastal Ocean Dynamics Experiment) 2000 off the coast of New Jersey (LEO-15 site)

measurements taken near local noon on 24 July 2000 at 39º 24.91' N, 74º, 11.78' W (station 19); cloudy sky, wind = 6 m/s

See Mobley et al, 2002, *Applied Optics* 41(6), 1035-1050 for details

Data Taken at the LEO-15 Site as Used to Model the In-Water Light Field^a Table 4.

"Most instruments have a nominal 10-nm bandwidth centered on the listed wavelengths.

HyCODE Data

ac-9, both filtered (CDOM absorption) and unfiltered (total *a* and *b*)

HydroScat-6 (*b^b*)

can get *B^p* from measured b_{b} *p*^{*b*_p}

can then use B_ρ to define a Fournier-Forand phase function with the same backscatter fraction (Mobley, 2002. *AO* 41(6), 1035-1050)

HyCODE Data

Also have VSF measurements (extremely rare) at 2 m depth at 530 nm from a novel Ukrainian instrument (Lee and Lewis, 2003. *J Atmos Ocean Tech* 20(4), 563-571)

HyCODE Data

Note that the measured B_p is much less than for the commonly used Petzold "average particle" phase function (0.0183), and B_n varies with depth and wavelength; value depends on type of particles: predominately phytoplankton near surface vs resuspended sediments near the bottom (18 m depth)

HyCODE Data: HydroLight vs E_d Measurements

measurements green: H with Petzold phase function

black:

red: H with FF phase function determined from measured *b^b /b*

blue: H with measured pf

HyCODE Data: HydroLight vs *L^u* Measurements

measurements green: H with Petzold phase function

black:

red: H with FF phase function determined from measured *b^b /b*

blue: H with measured pf

HyCODE Data: HydroLight vs *L^u /E^d* Measurements

black: measured by Hyper-TSRB (Satlantic)

purple dots: measured by OCP (Ocean Color Profiler; Satlantic)

green: H with Petzold phase func.

red: H with FF pf determined from measured *b^b /b*

blue: H with measured pf

HyCODE Data: HydroLight vs K_{Lu} Measurements

Fig. 16. Comparison of measured and HYDROLIGHT-predicted diffuse attenuation for upwelling radiance at 591 nm. Solid curve, measured values; dotted curve, predicted values with the FF phase function; dashed curve, predicted values with the measured phase function; dashed-dotted curve, predicted values with the Petzold phase function.

The Mobley et al. 2002 paper showed that the exact shape of the phase function makes only a few per cent difference in *E*d , *L*^u , *R*rs, etc, so long as the backscatter fraction B_{p} $b_{\rm bp}/b_{\rm p}$ is correct

Tzortziou et al, "Bio-optics of the Chesapeake Bay from measurements and radiative transfer closure." *Estuarine, Coastal and Shelf Science* (2006).

She shows how to "do it right" in taking and processing data, and modeling it with HydroLight.

Complex case 2 water

Read this paper!

Fig. 4. Particulate absorption (sum of absorption by phytoplankton and nonalgal particles) in the 290-750 nm wavelength region, measured spectrophotometrically for station PI, on 28 September 2001. The residual non-zero particulate absorption at 715 nm is shown more clearly in the inset figure.

Can't use the simple ac-9 scattering correction that assumes *a*^p (715) = 0

Improved ac-9 scattering correction ($a_{\text{p}}(715) \neq 0$) gives better *L*_u and *E*_d

Improved ac-9 scattering correction and phase function via $b_{\rm b}/b$ and fluorescence gives better $L_{\rm w}$

Table 2

Improvement of agreement between measured and model-estimated L_w as information on the specific IOPs measured at station PI (28 September 2001) is successively incorporated into the model. The final agreement between data and model demonstrates the good optical closure obtained at this study site after applying the results from our detailed measurements to properly account in the radiative transfer modeling for the observed optical characteristics

Fig. 3. L_w spectra estimated using: (1) a Petzold "average particle" scattering phase function (stars), (2) an FF scattering phase function with a constant backscattering ratio, $b_b/b = 0.015$ (filled circles); and (2) an FF scattering phase function as determined by measured wavelength- and depth-dependent b_b/b (filled squares). Measured L_w are shown as open diamonds. Percent differences in L_w between measurements and model estimations are shown in the inset figure (percent differences estimated as $(L_{w (model)} - L_{w (data)})/L_{w (data)}$).

Phase function effects on L_w

*L*_u and *E*_d comparisons for 3 stations and 3 wavelengths

Table 3

Percent differences in measured and model-estimated $L₀$ and E_d (443, 554, 670 nm) at 1 m depth. Comparisons are shown for measurements representative of the most turbid waters we observed in the Bay (22 May 2002). Percent differences were estimated as:

$$
\frac{L_{u \text{(model)}} - L_{u \text{(data)}}}{\frac{1}{2}(L_{u \text{(model)}} + L_{u \text{(data)}})} 100 \text{ (and similarly for } E_d)
$$

Comparison of all *L*_u and *E*_d measurements

Fig. 8. Comparison between model-estimated and in-situ measured (a) $E_d(z)$ (in μ W nm⁻¹ cm⁻²) and (b) $L_u(z)$ values (in μ W nm⁻¹ cm⁻² sr⁻¹) at depths 0-6 m, for all cruises-stations that comparisons with the radiative transfer model were performed. Comparisons within the first 3 m are shown as dark circles ($R^2 = 0.99$ for E_d , $R^2 = 0.95$ for L_a), while comparisons for depths below 3 m are shown as open circles ($R^2 = 0.95$ for E_d , $R^2 = 0.92$ for L_a) (the 1:1 line is also shown for comparison).

 L_w and HydroLight comparisons at various stations Note: errors bars are std dev from 3 measurments

wavelength (nm)

L_{w} and HydroLight comparisons at various stations

Table 4

Percent differences in estimated L_w values (412-685 nm) using in-situ measurements and model simulations. Percent differences were estimated as:

 $\frac{L_\mathrm{w(model)}-L_\mathrm{w(data)}}{ \frac{1}{2}\big(L_\mathrm{w(model)}+L_\mathrm{w(data)}\big)}100$

What else could be added to this study?

Measured vs HydroLight for CICORE Station ER01 CICORE data and analysis by Heidi Dierssen, Univ. Conn.; used measured ac-9 *a* and *b*; best-guess Fournier-Forand phase function, etc.]

3 instruments & 2 HydroLight

Note that the 3 instruments disagree by about the same amount as the two H simulations (using different guesses for the phase function)

You Get the Idea

You do the best you can with the data you have. Sometimes very good, sometimes not so good, sometimes completely useless. That's science.

If you didn't measure the VSF, can you get the backscatter fraction from b $_{\rm b}$ /b? If not, treat b $_{\rm b}$ /b as a "fitting parameter" and tweak to get the best fit for *R_{rs}*, for example.

Even if you can't get agreement between measured and modeled E_d and L_u, for example, can you get agreement with L_u/E_d or with K_d?

Compare as many things as possible, e.g., the measured E_d from the HyperPro and from the ship deck cell and with H's default sky irrad model.

The disagreements are often where you learn the most.

Play around with HydroLight. Have fun!

