

# 2015 Summer Course on Optical Oceanography and Ocean Color Remote Sensing

## HydroLight Model-Data Closure

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University of Maine  
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# Overview

- Show two examples of output from an advanced simulation that show what is necessary to achieve model-data closure, i.e., getting all of your inputs to H and outputs from H to agree with your measurements
- Lab 2 will demonstrate how to build up IOPs from several different components (water + phytoplankton + CDOM + ....), input your own data (ac9, bottom reflectance, etc) using HydroLight Standard Formats, etc.

# Measurements Necessary for Model-Data Closure

## HydroLight inputs

- absorption coef  $a(z,\lambda)$  (e.g., from ac-9 or spectrophotometer)
- scattering coef  $b(z,\lambda)$  (e.g., from ac-9)
- scattering phase function  $\beta(z,\lambda,\psi)$  (almost never measured, but may have backscatter fraction  $B = b_b/b$  from  $b_b$  (e.g., HydroScat or EcoVSF) and  $b$  (ac-9))
- boundary conditions: sea state (wind speed); sun location and sky conditions (usually model), bottom reflectance (in shallow water)

## HydroLight outputs

- radiometric variables (radiances and irradiances; usually measure  $L_u(z,\lambda)$  and  $E_d(z,\lambda)$  at a minimum)
- apparent optical properties ( $K_d$ ,  $R$ ,  $R_{rs}$  etc. obtained from radiometric measurements). The most common for remote sensing is remote sensing reflectance  $R_{rs}$  (measure  $E_d(\text{air})$  and  $L_u(z)$  and extrapolate upward from underwater  $L_u$ , or estimate  $R_{rs}$  using above-surface techniques)

# Comprehensive Data Sets Are Extremely Scarce

Data set from ONR HyCODE (Hyperspectral Coastal Ocean Dynamics Experiment) 2000 off the coast of New Jersey (LEO-15 site)

measurements taken near local noon on 24 July 2000 at 39° 24.91' N, 74°, 11.78' W (station 19); cloudy sky, wind = 6 m/s

See Mobley et al, 2002, *Applied Optics* 41(6), 1035-1050 for details

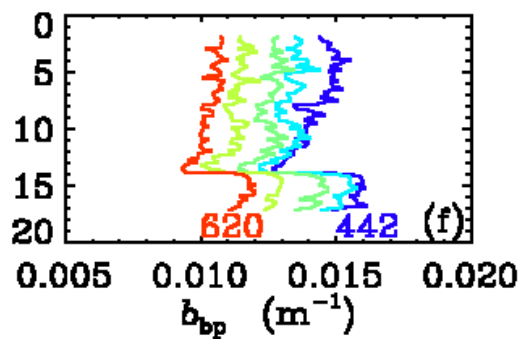
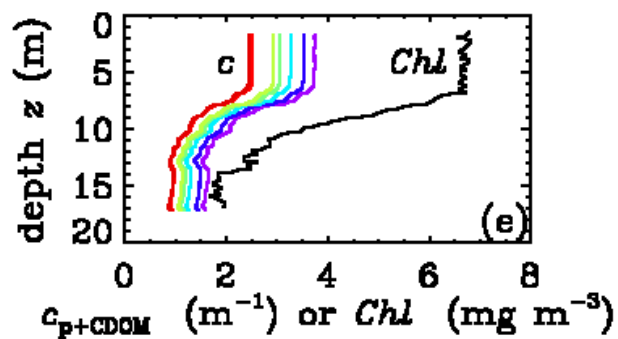
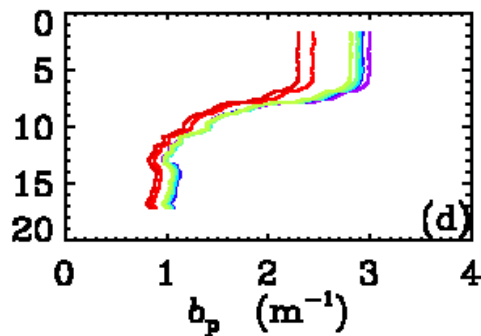
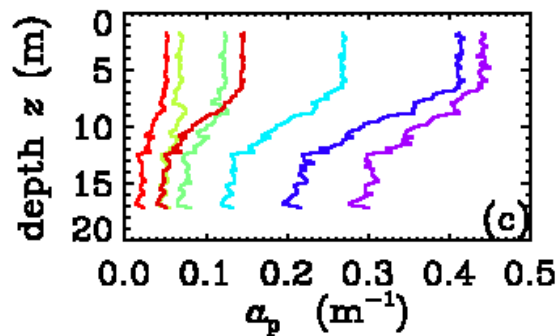
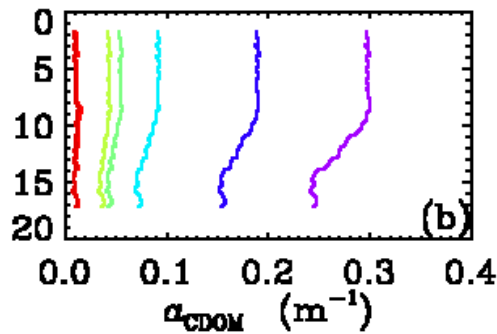
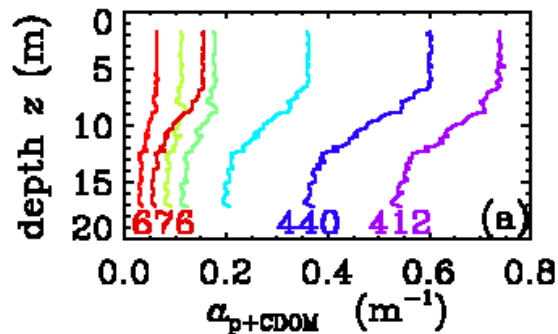
Table 4. Data Taken at the LEO-15 Site as Used to Model the In-Water Light Field<sup>a</sup>

Quantity Measured	Instrument	Nominal Wavelength (nm)
Total $a(z, \lambda)$ , total $c(z, \lambda)$	Unfiltered ac-9	412, 440, 488, 532, 555, 650, 676, 715
Dissolved $a(z, \lambda)$	Filtered ac-9	412, 440, 488, 532, 555, 650, 676, 715
Backscatter $b(z, \lambda)$ derived from VSF at $\psi = 140$ deg	HydroScat-6	442, 488, 532, 555, 620
Backscatter $b(z, \lambda)$ derived from VSF at $\psi = 100, 125,$ and $150$ deg	ECO-VSF	530
VSF ( $\psi = 0.6$ – $179.6$ deg)	VSM	530
$E_d(z, \lambda)$ and $L_u(z, \lambda)$	OCP	412, 443, 489, 533, 555, 591, 683
Sky $E_d(\lambda)$	Multichannel visible detector system	412, 443, 489, 533, 555, 591, 683
Sky $E_d(\lambda)$ and $L_u(z = 0.6 \text{ m}, \lambda)$	Hyper-TSRB	123 wavelengths between 396 and 798

<sup>a</sup>Most instruments have a nominal 10-nm bandwidth centered on the listed wavelengths.

Think about what should be measured in a field experiment, but never is (cost, lack of interest, ignorance, politics, ...)

# HyCODE Data



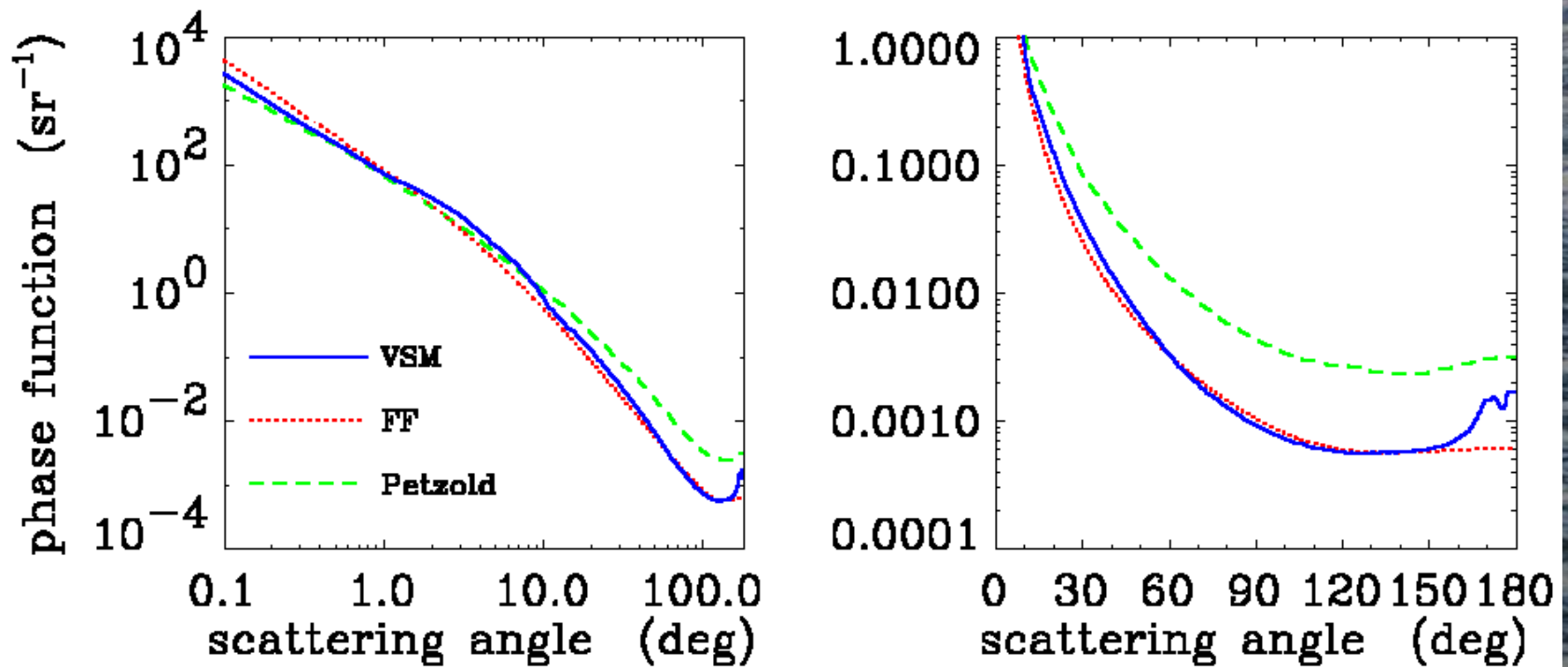
ac-9, both filtered (CDOM absorption) and unfiltered (total  $a$  and  $b$ )

HydroScat-6 ( $b_p$ )

can get  $B_p$  from measured  $b_{bp}/b_p$

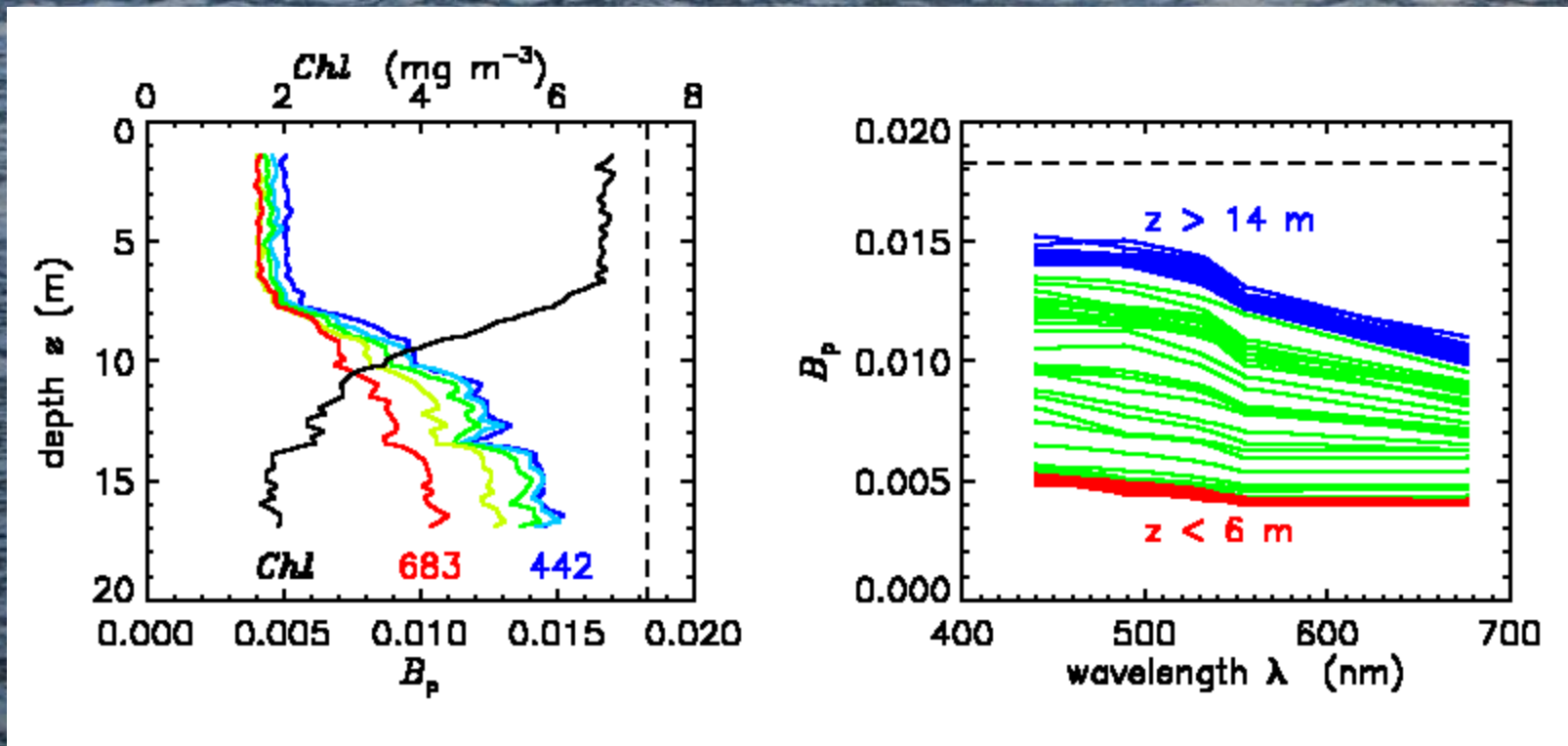
can then use  $B_p$  to define a Fournier-Forand phase function with the same backscatter fraction (Mobley, 2002. AO 41(6), 1035-1050)

# HyCODE Data



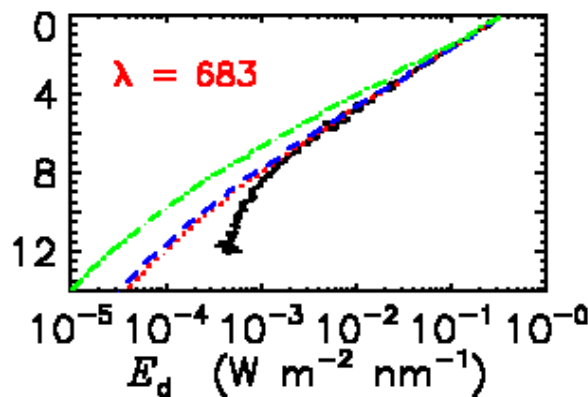
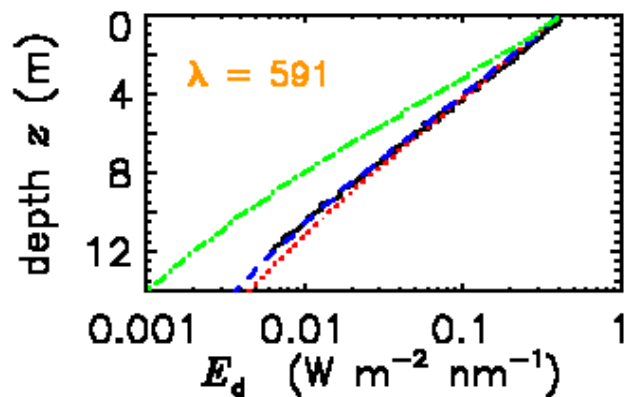
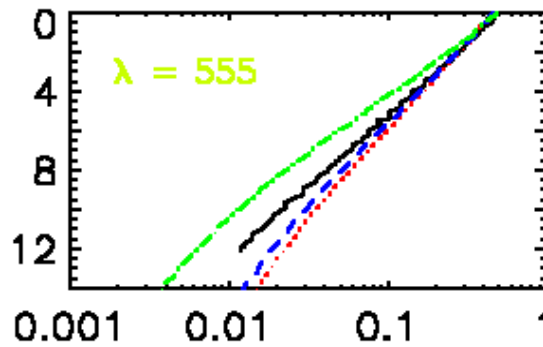
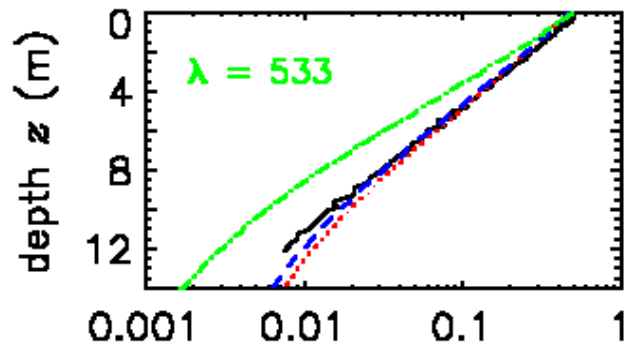
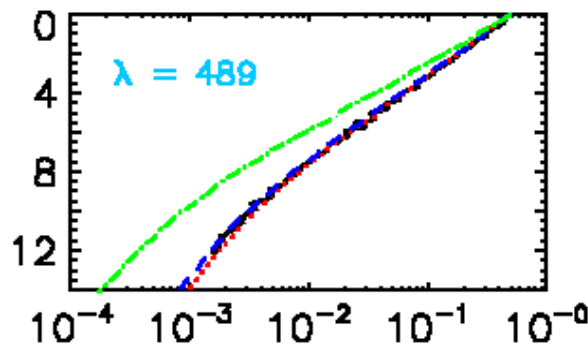
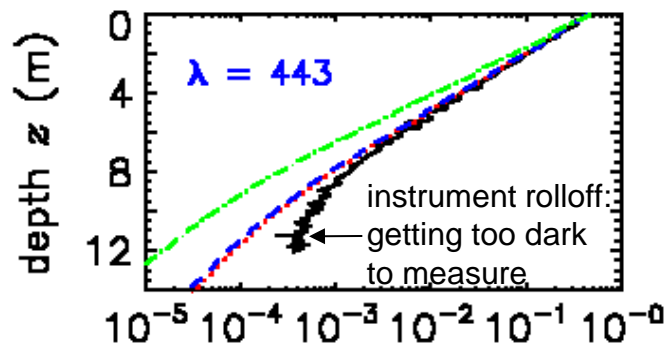
Also have VSF measurements (extremely rare) at 2 m depth at 530 nm from a novel Ukrainian instrument (Lee and Lewis, 2003. *J Atmos Ocean Tech* 20(4), 563-571)

# HyCODE Data



Note that the measured  $B_p$  is much less than for the commonly used Petzold “average particle” phase function (0.0183), and  $B_p$  varies with depth and wavelength; value depends on type of particles: predominately phytoplankton near surface vs resuspended sediments near the bottom (18 m depth)

# HyCODE Data: HydroLight vs $E_d$ Measurements



black:  
measurements

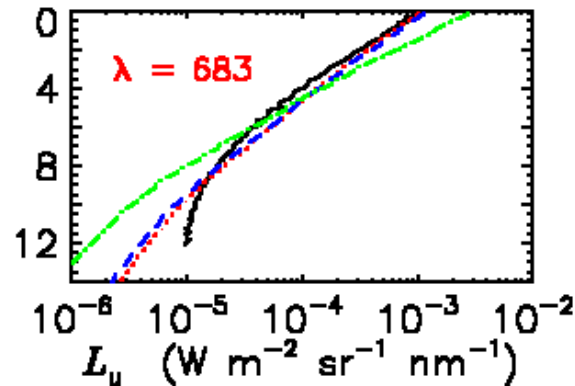
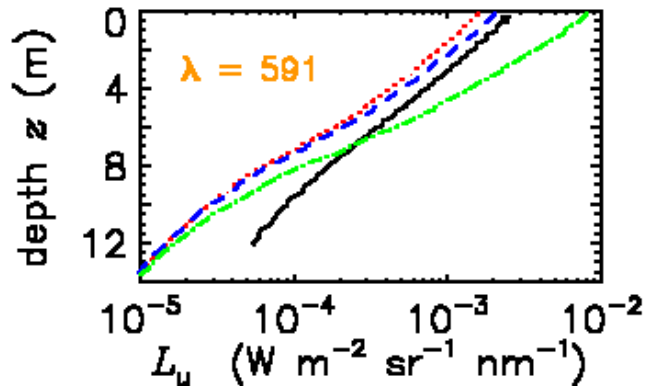
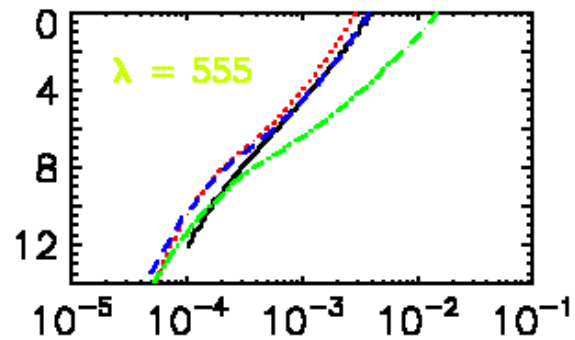
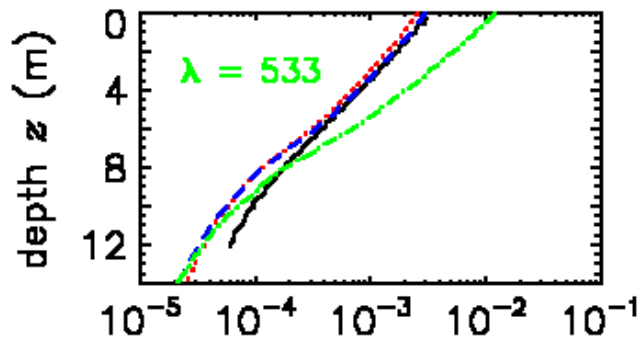
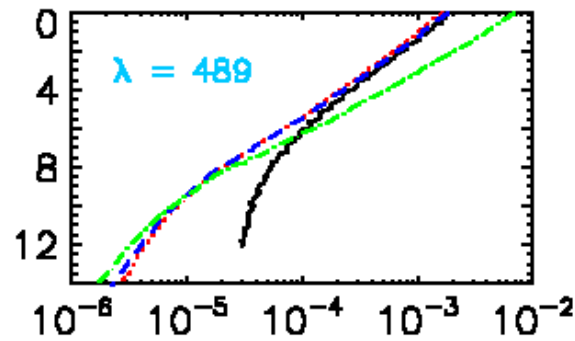
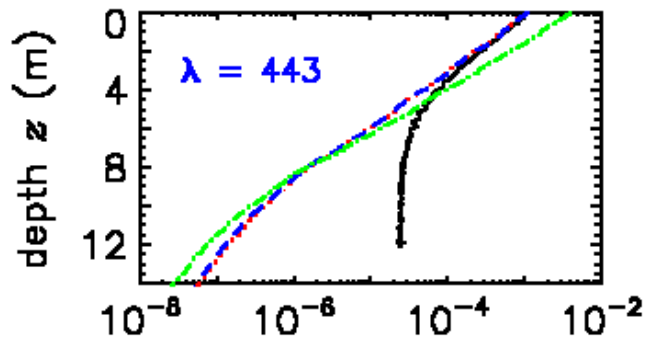
green: H with  
Petzold phase  
function

red: H with FF  
phase function  
determined from  
measured  $b_p/b$

blue: H with  
measured pf



# HyCODE Data: HydroLight vs $L_u$ Measurements



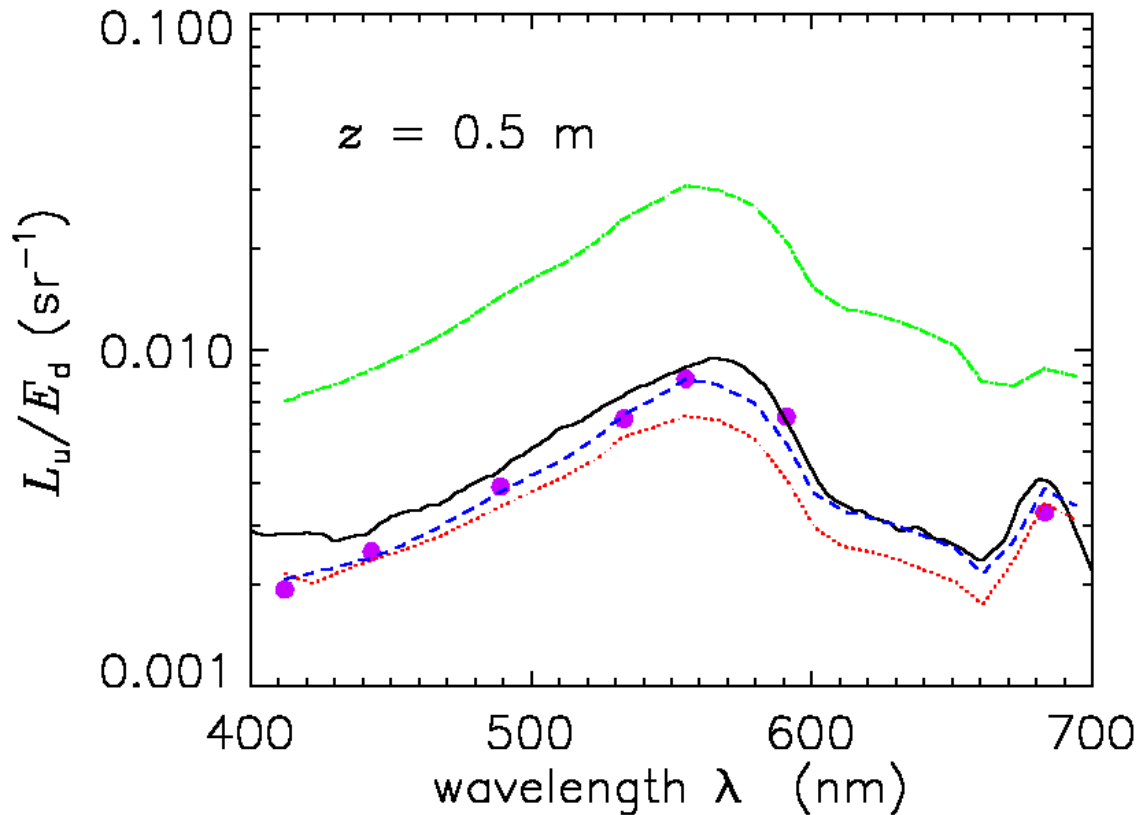
black:  
measurements

green: H with  
Petzold phase  
function

red: H with FF  
phase function  
determined from  
measured  $b_p/b$

blue: H with  
measured pf

# HyCODE Data: HydroLight vs $L_u/E_d$ Measurements



black: measured by  
Hyper-TSRB  
(Satlantic)

purple dots:  
measured by OCP  
(Ocean Color  
Profiler; Satlantic)

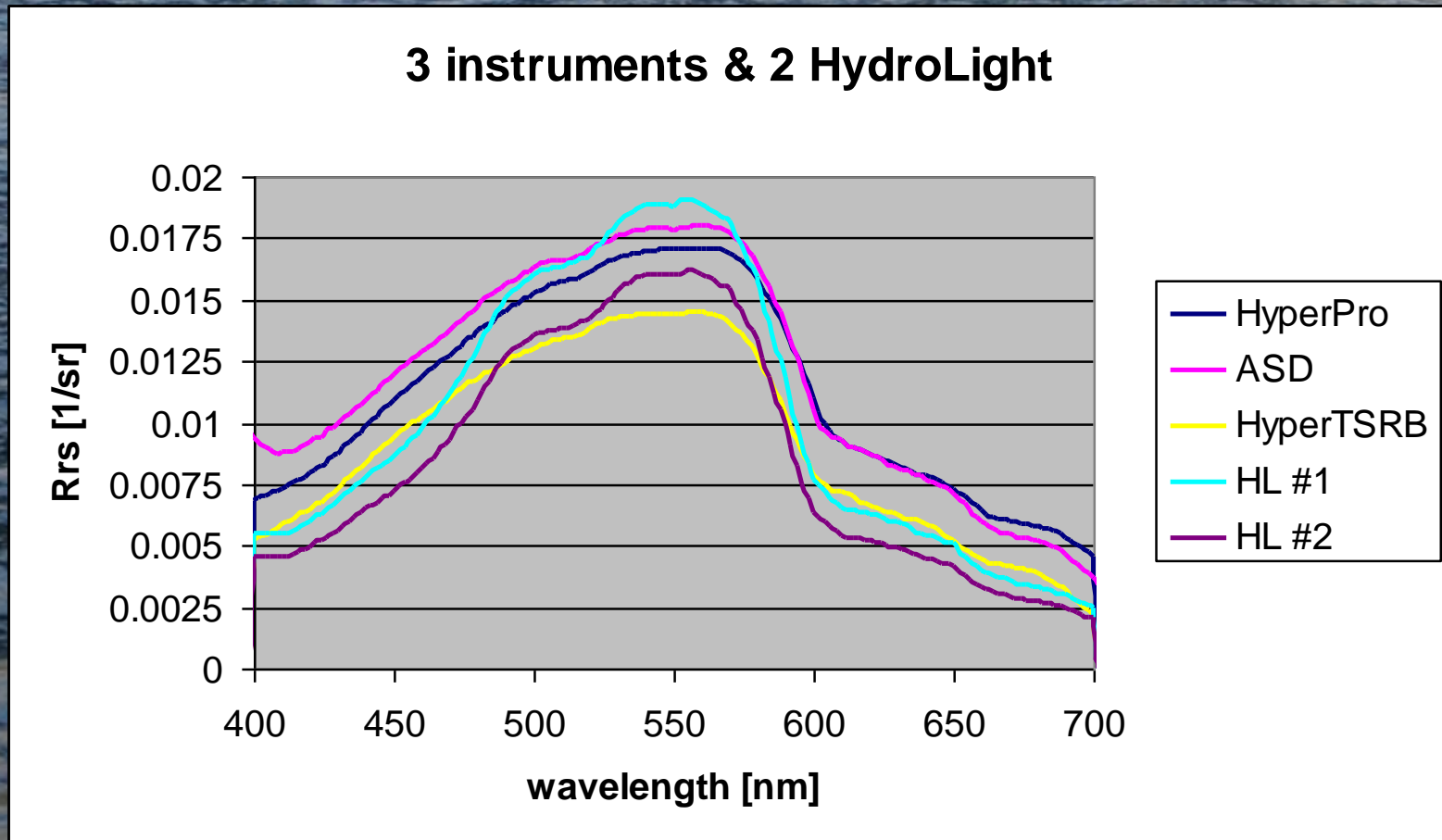
green: H with  
Petzold phase func.

red: H with FF pf  
determined from  
measured  $b_p/b$

blue: H with  
measured pf

# Measured vs HydroLight for CICORE Station ER01

CICORE data and analysis by Heidi Dierssen, Univ. Conn.; used measured ac-9 *a* and *b*; best-guess Fournier-Forand phase function, etc.]



Note that the 3 instruments disagree by about the same amount as the two H simulations (using different guesses for the phase function)

Tzortziou et al, *Estuarine & Coastal Syst. Sci.* (2006) show how to “do it right” in taking and processing data, and modeling it with HydroLight. **Read this paper!**



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## Bio-optics of the Chesapeake Bay from measurements and radiative transfer closure

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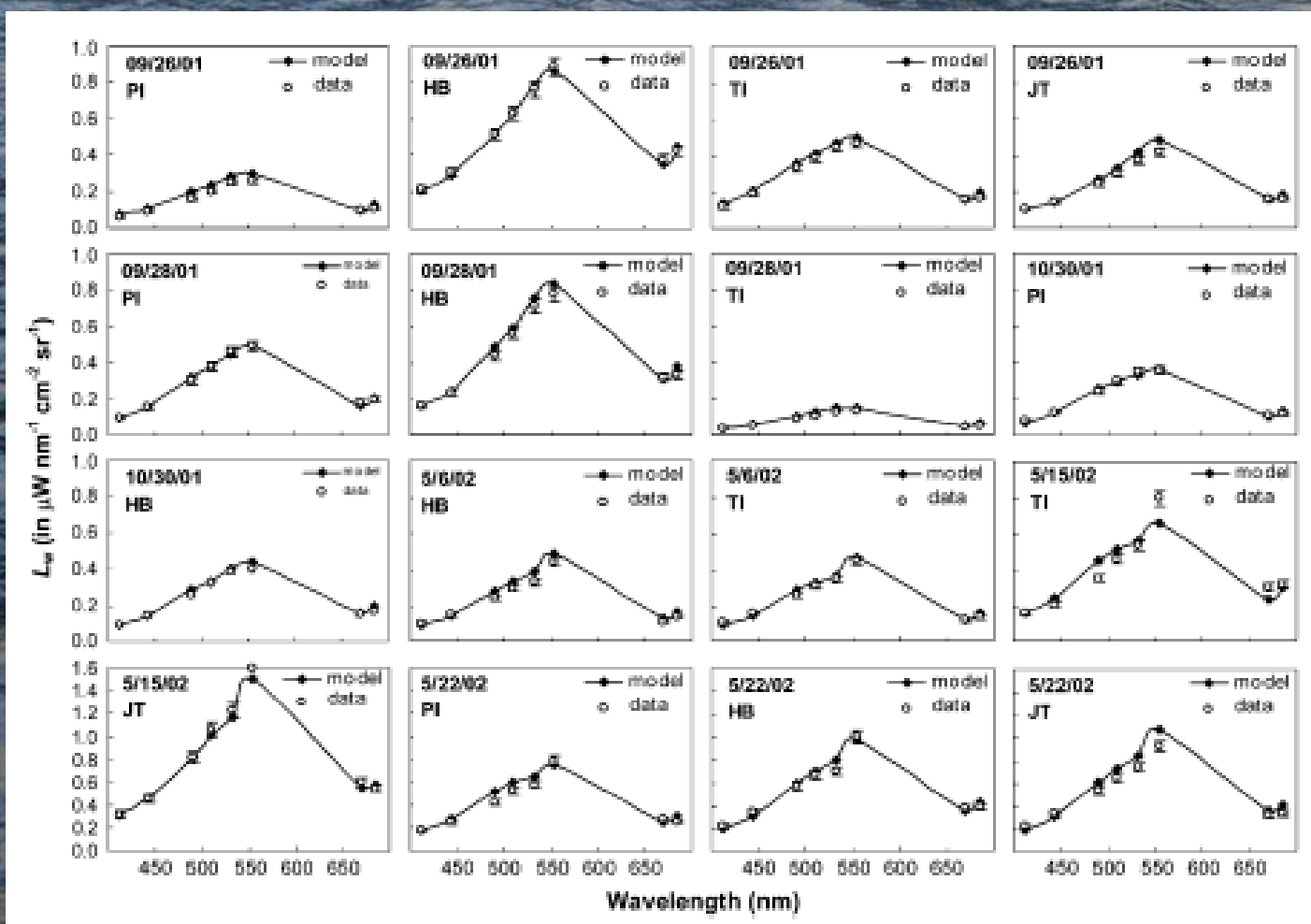
### Abstract

We combined detailed bio-optical measurements and radiative transfer modeling to perform an ‘optical closure’ experiment for an optically complex and biologically productive region of the Chesapeake Bay. We used this experiment to evaluate certain assumptions commonly used in bio-optical models, and to investigate which optical characteristics are most important to accurately model and interpret remote sensing ocean-color observations in these Case 2 waters. Direct measurements were made of the magnitude, variability, and spectral characteristics of backscattering and absorption that are critical for accurate parameterizations in satellite bio-optical algorithms and underwater radiative transfer simulations. We found that the ratio of backscattering to total scattering (i.e. the backscattering fraction,  $b_b/b$ ) varied considerably depending on particulate loading, distance from land, and mixing processes, and had an average value of 0.0128 at 530 nm. Incorporating information on the magnitude, variability, and spectral characteristics of particulate backscattering into the radiative transfer model, rather than using a volume scattering function commonly assumed for turbid waters, was critical to obtaining agreement between model calculations and measured radiometric quantities. In-situ measurements of absorption coefficients need to be corrected for systematic overestimation due to scattering errors, and this correction commonly employs the assumption that absorption by particulate matter at near-infrared wavelengths is zero. Direct measure-

# Measured vs HydroLight for Chesapeake Bay

Case 2 water. From Tzortziou et al, Estuarine & Coastal Syst. Sci. (2006).

$L_w$  ( $\mu\text{W cm}^{-2} \text{sr}^{-1} \text{nm}^{-1}$ )



wavelength (nm)

# Measured vs HydroLight for Chesapeake Bay

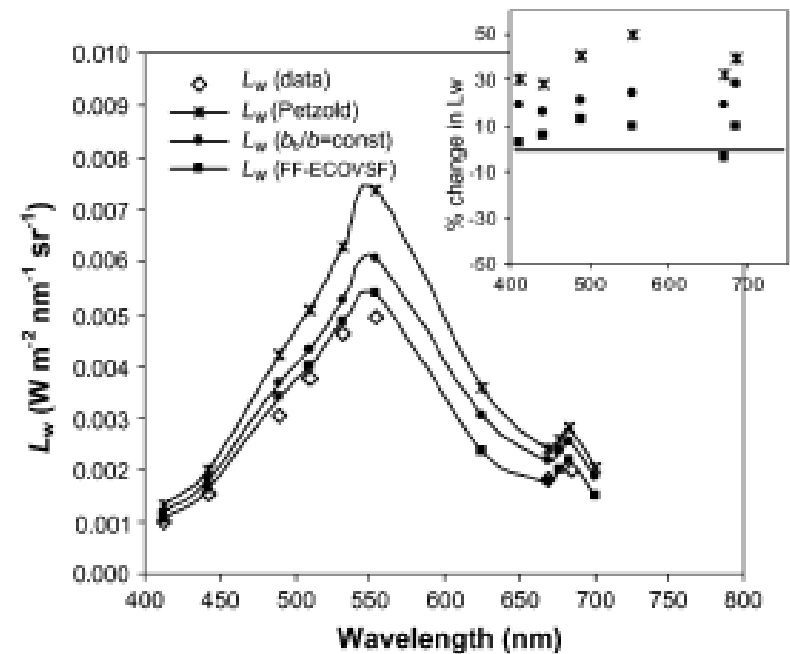
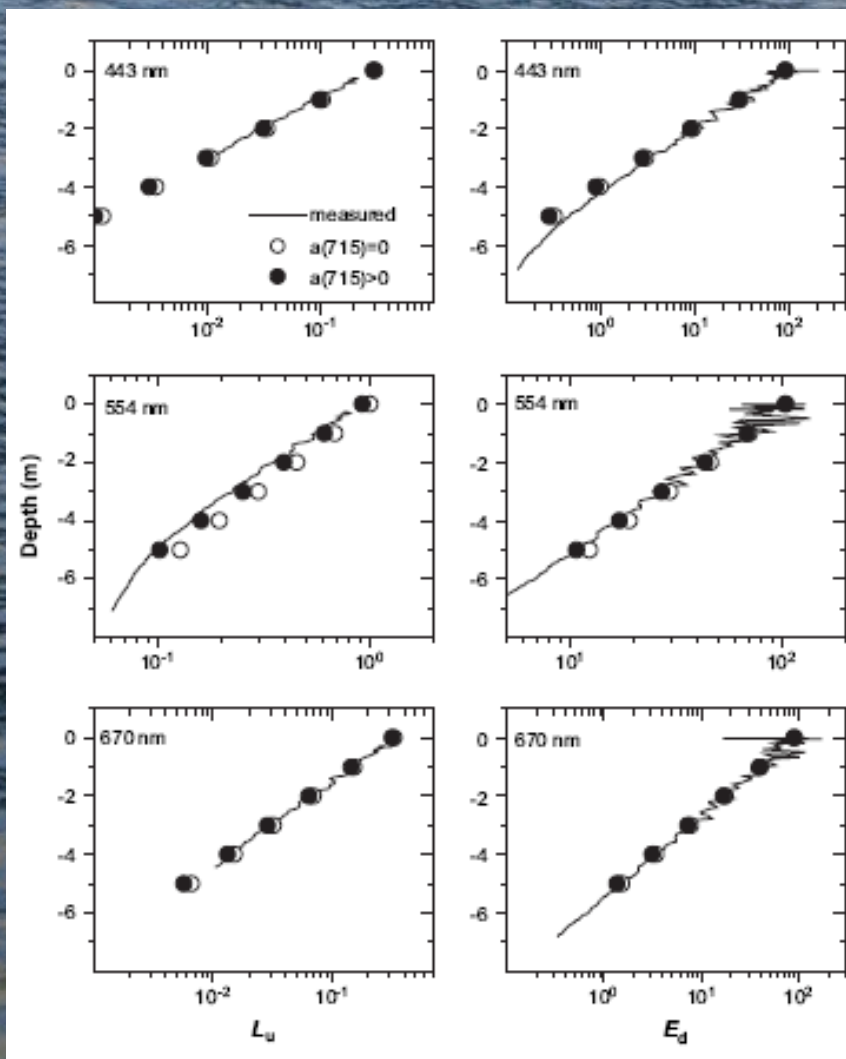


Fig. 3.  $L_w$  spectra estimated using: (1) a Petzold "average particle" scattering phase function (stars), (2) an FF scattering phase function with a constant backscattering ratio,  $b_p/b = 0.015$  (filled circles); and (3) an FF scattering phase function as determined by measured wavelength- and depth-dependent  $b_p/b$  (filled squares). Measured  $L_w$  are shown as open diamonds. Percent differences in  $L_w$  between measurements and model estimations are shown in the inset figure (percent differences estimated as  $(L_{w(model)} - L_{w(data)})/L_{w(data)}$ ).

other examples from  
Tzortziou et al. 2006

## You Get the Idea

You do the best you can with the data you have. Sometimes very good, sometimes not so good, sometimes completely useless. That's science.

If you didn't measure the VSF, can you get the backscatter fraction from  $b_b/b$ ? If not, treat  $b_b/b$  as a "fitting parameter" and tweak to get the best fit for  $R_{rs}$ , for example.

Even if you can't get agreement between measured and modeled  $E_d$  and  $L_u$ , for example, can you get agreement with  $L_u/E_d$  or with  $K_d$ ?

Compare as many things as possible, e.g., the measured  $E_d$  from the HyperPro and from the ship deck cell and with H's default sky irradiance model.

The disagreements are often where you learn the most.

Play around with HydroLight. Have fun!

# Following Marco Polo Across the Silk Road in Western China

Photo by Curt Mobley

